

Sediment trap time series from the North Pacific Ocean

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Logistical support for
western Pacific traps

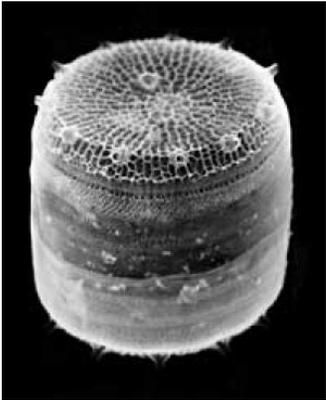
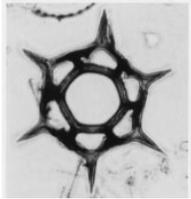
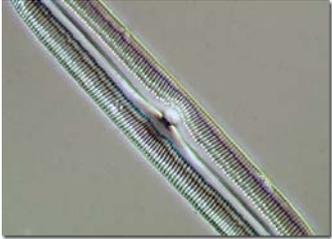
Officers and Crew
of the *JP Tully*

TIME SERIES FROM:

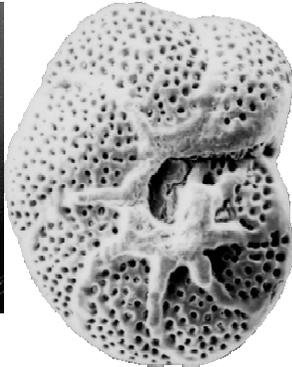
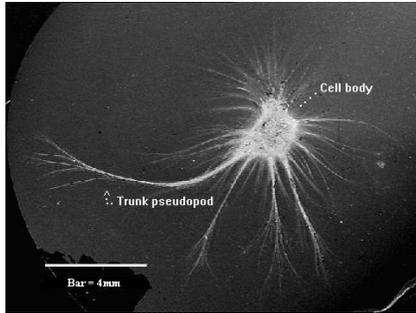
- Ocean Station PAPA
- Alaska Gyre and North Pacific Western Gyre
- Line P with emphasis at La Perouse Bank

Marine sediments are composed of...

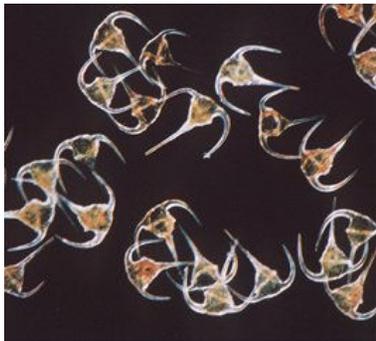
biogenic silica



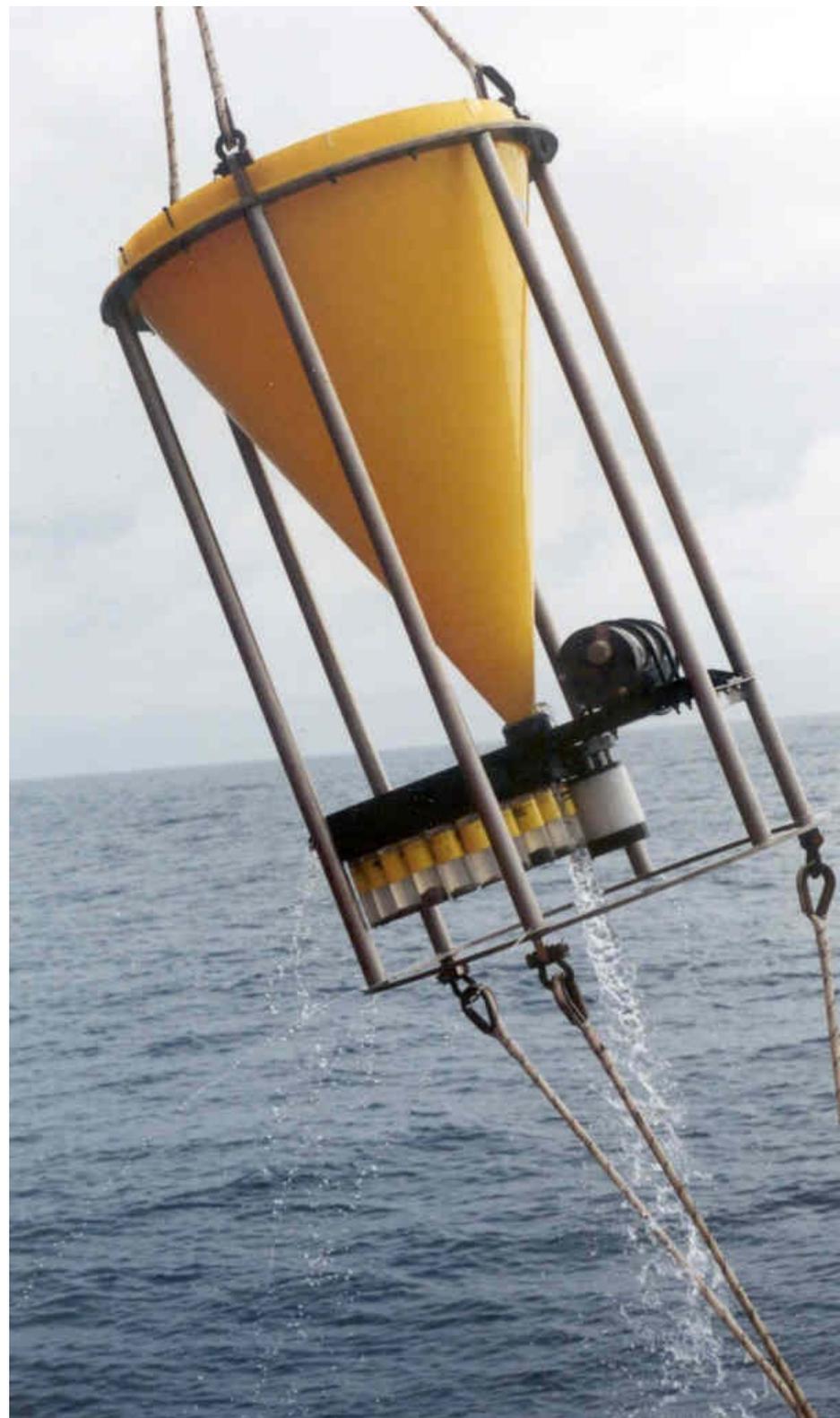
calcium carbonate



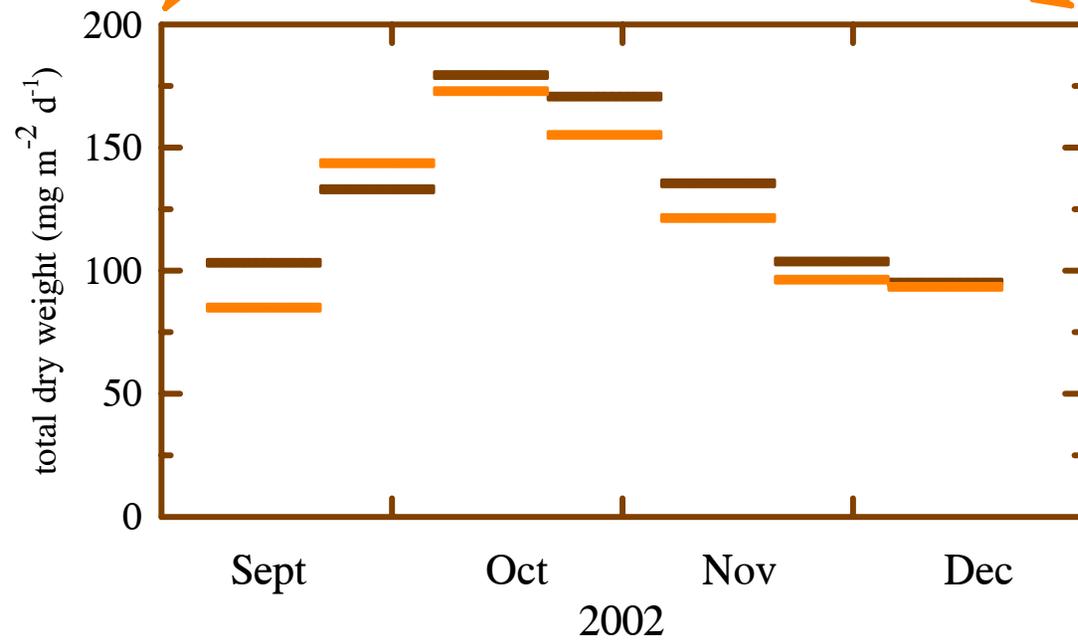
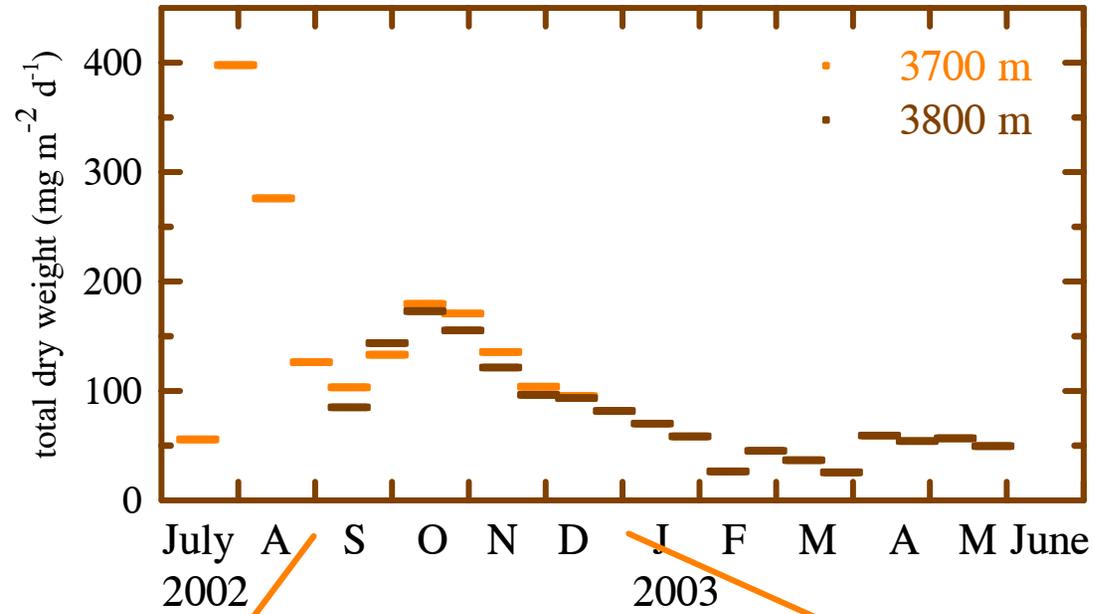
organic matter



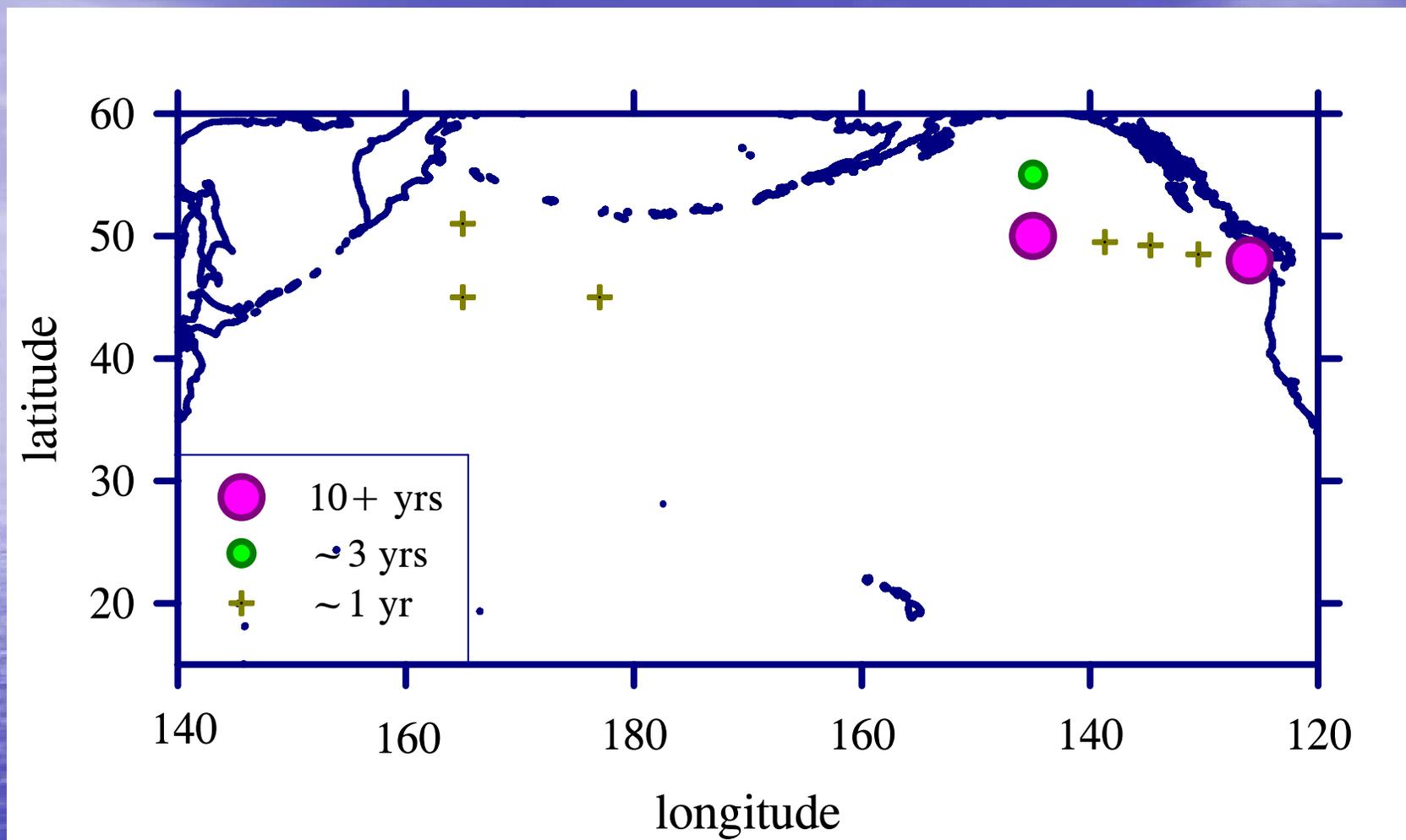
and rocks!



OSP: 2002-2003
overlapping deployments
3700 m - 3800 m



Sediment-trap moorings deployed by IOS



OSP: 1982 to present (3 depths)

3 stations in the Western Pacific:
August, 1991 to June, 1992

La Perouse Bank: 1986 to present

3 stations along line P:

Alaska Gyre: 1990 to 1992

sporadically from 1995 to 1997

TIME SERIES FROM:

➤ Ocean Station PAPA

➤ Alaska Gyre and North Pacific Western Gyre

➤ Line P with emphasis at La Perouse Bank

200 m

1 year of fluxes from OSP

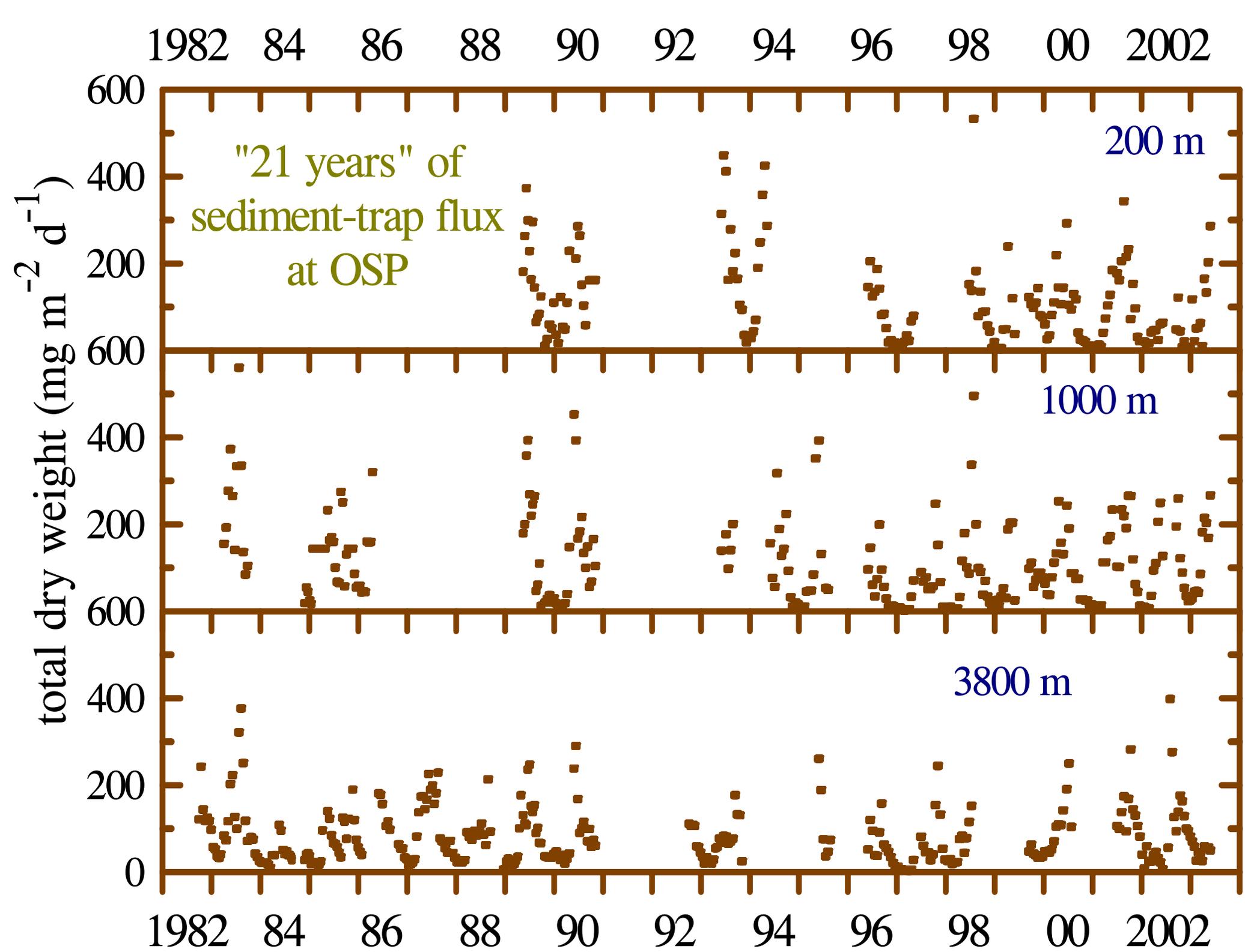


1000 m



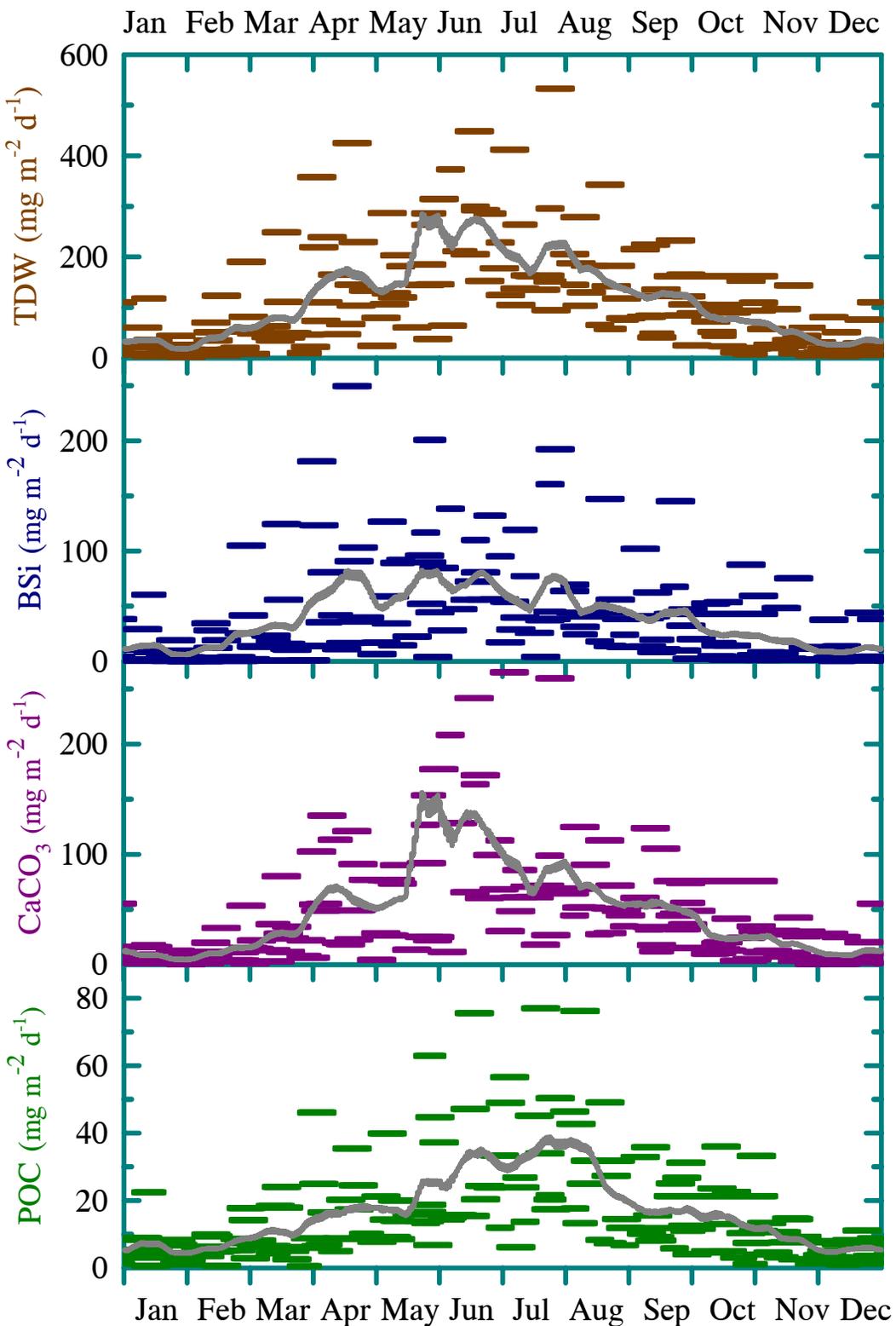
3800 m



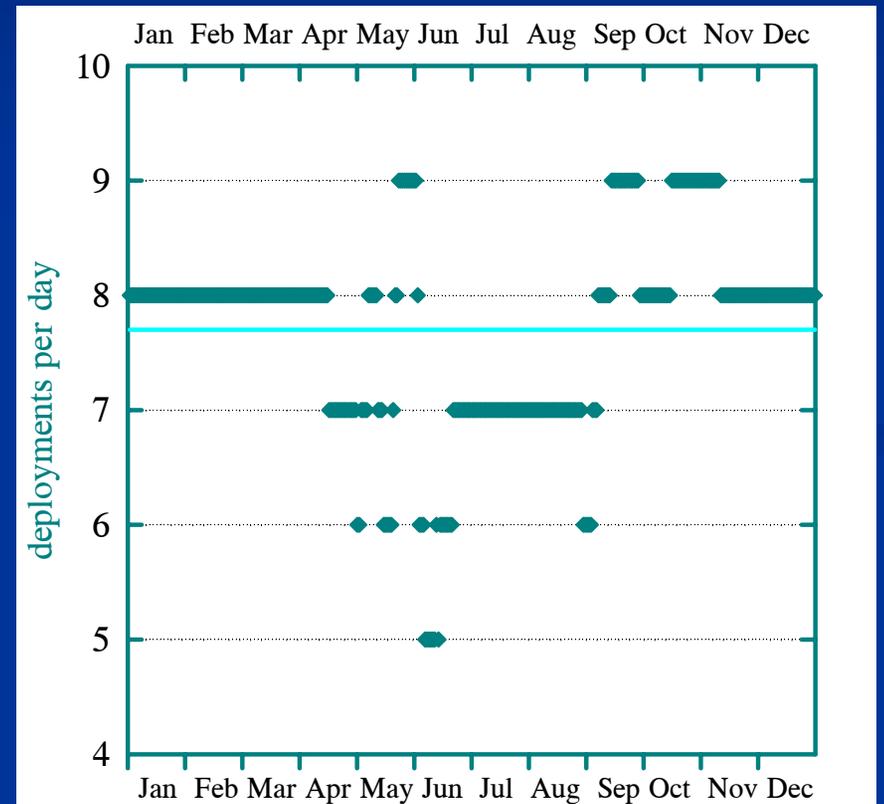


characterising seasonality and “filling gaps”
to generate a single time-series

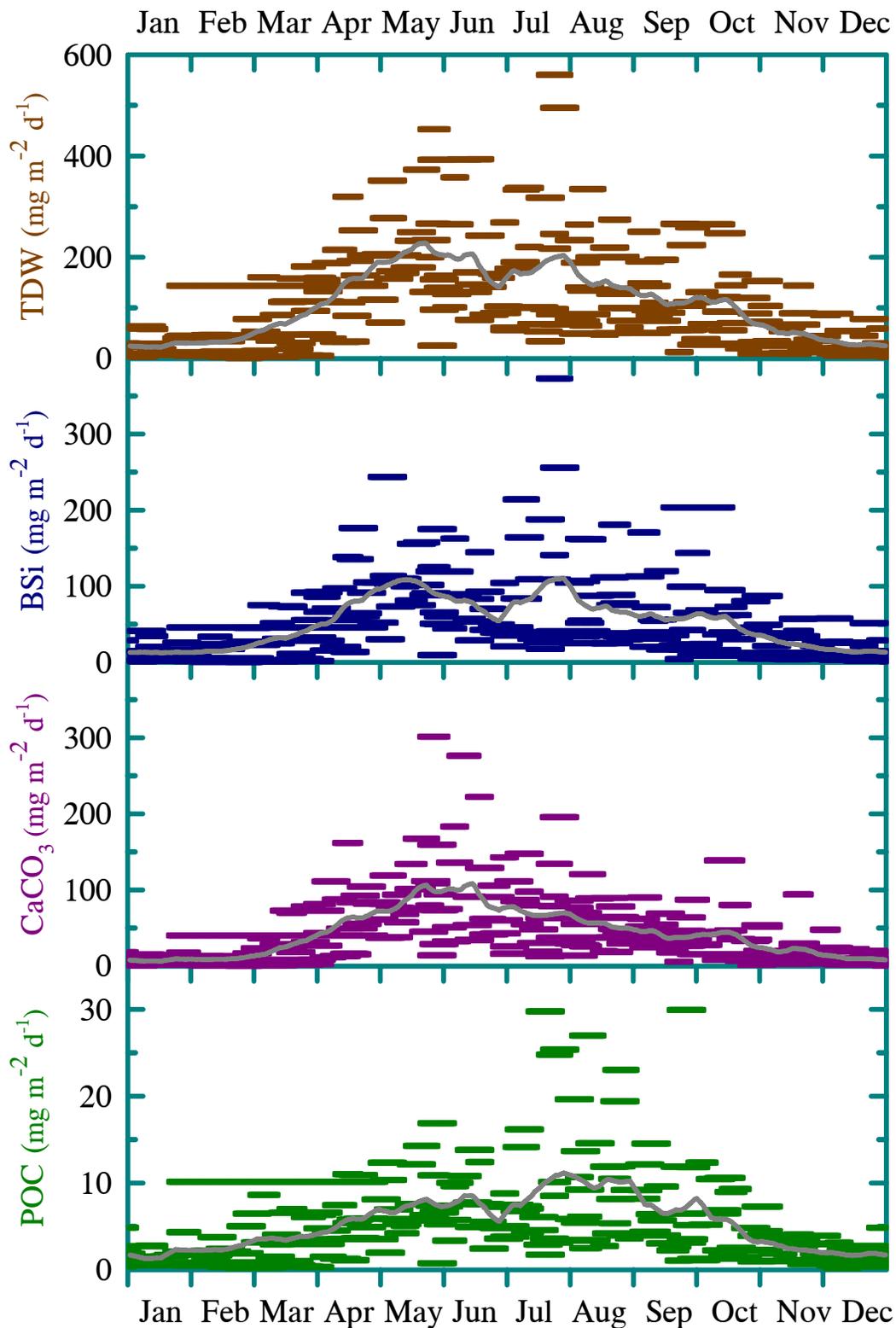




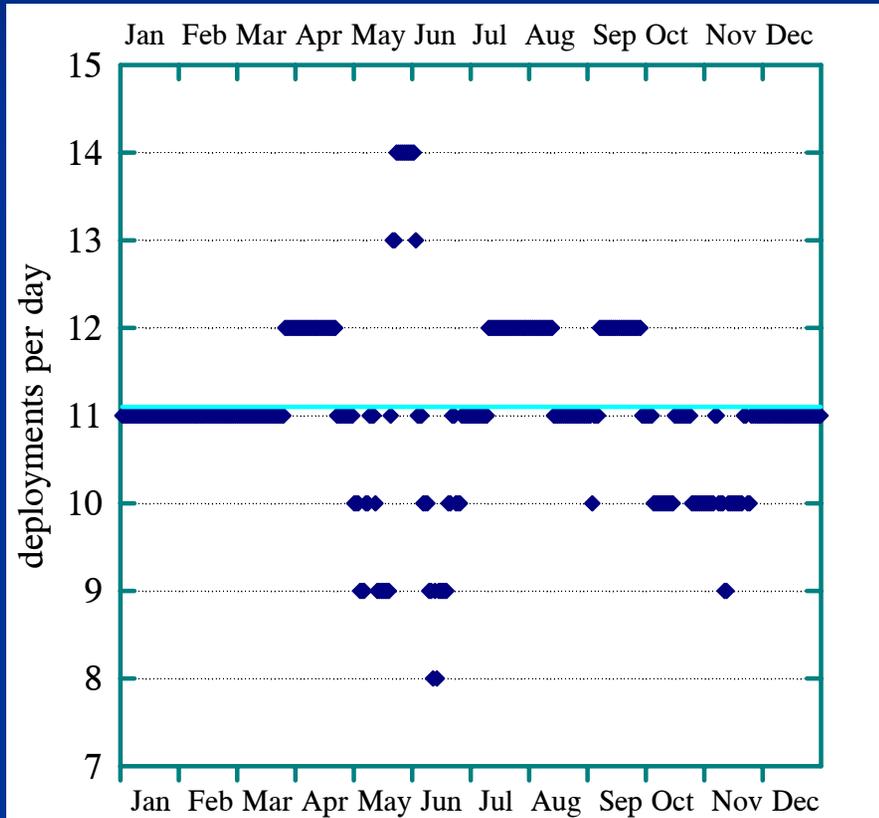
200 m fluxes at OSP sorted by Julian day



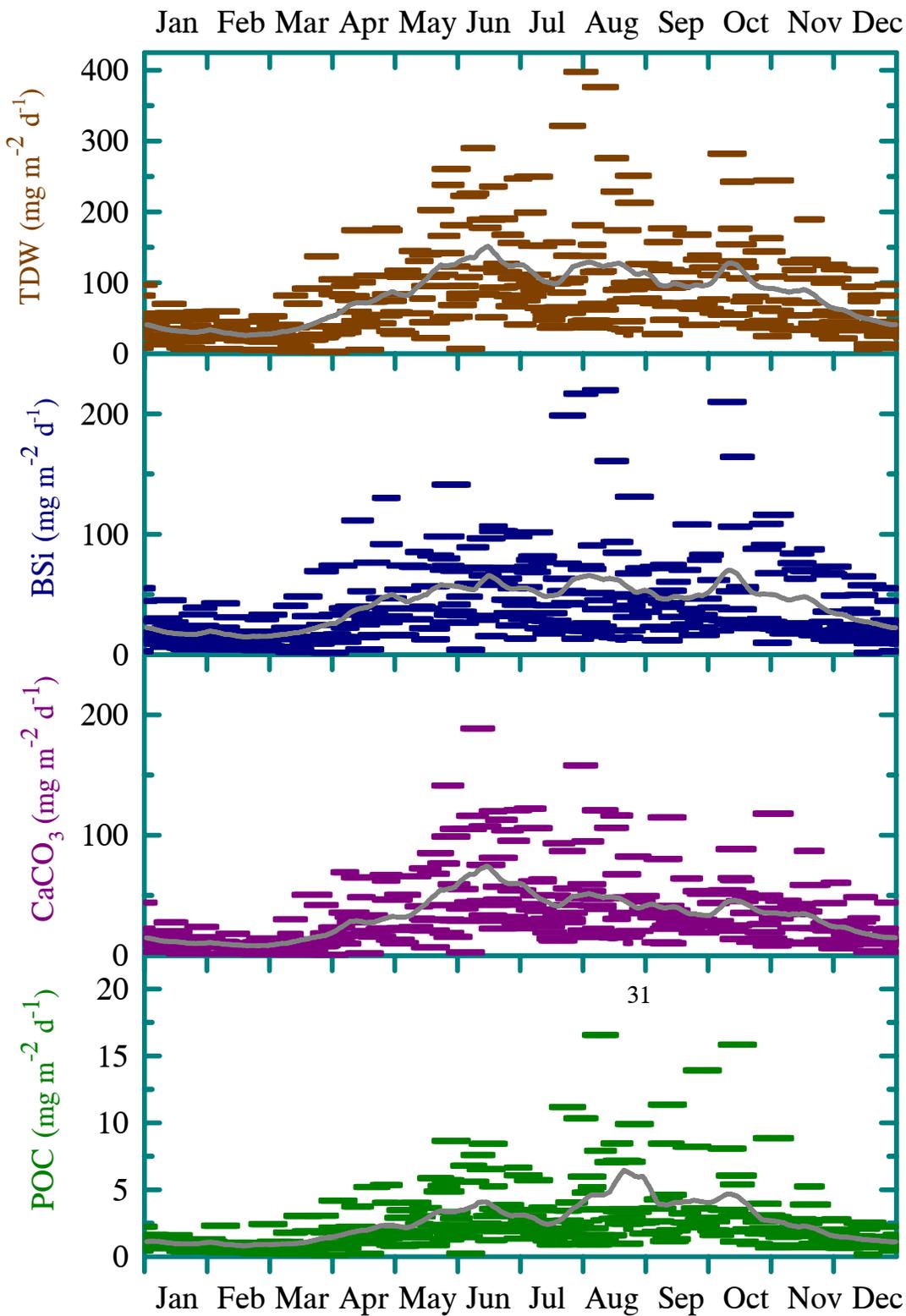
7.7 years of data



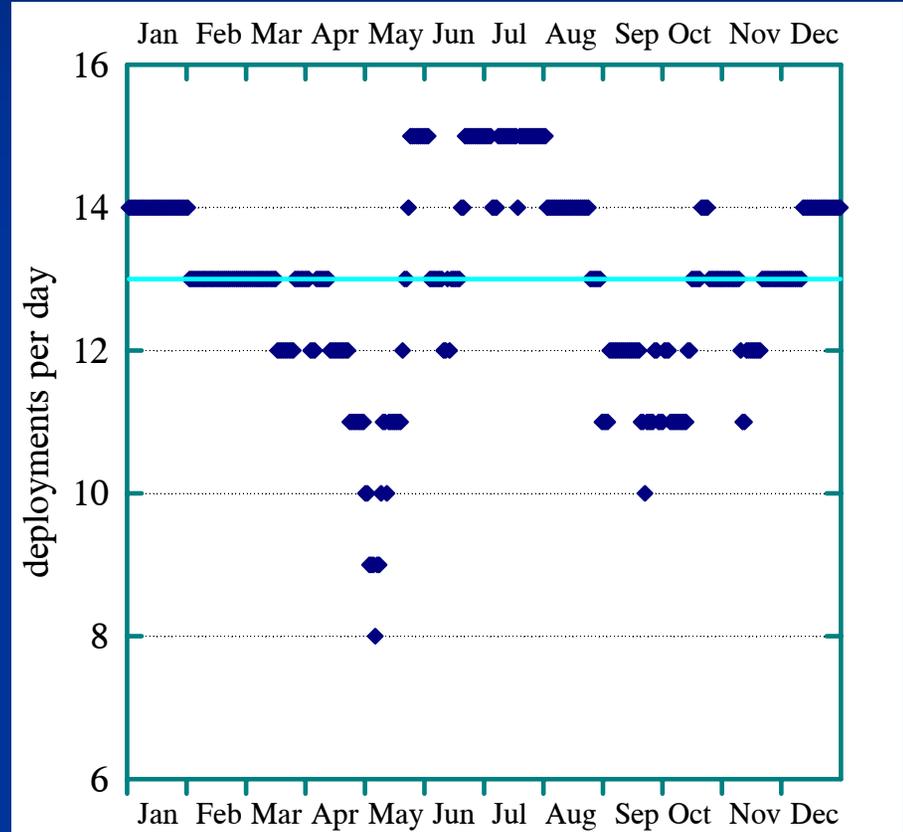
1000 m fluxes at OSP sorted by Julian day



11 years of data



3800 m fluxes at OSP sorted by Julian day



13 years of data

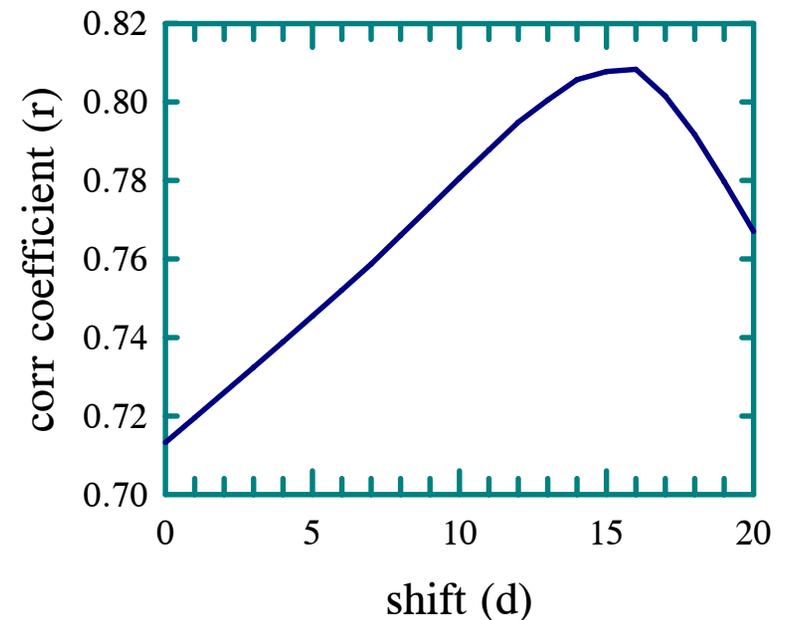
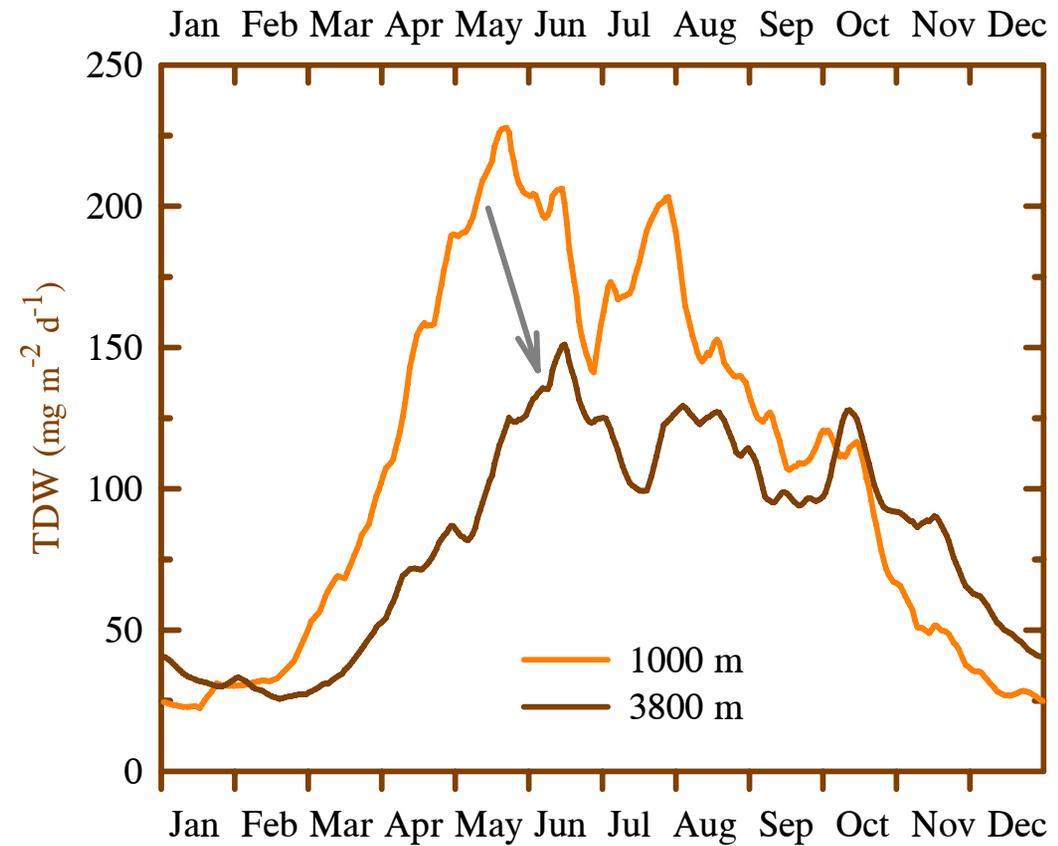
TIME-LAG IN FLUX BETWEEN SHALLOW AND DEEP TRAPS

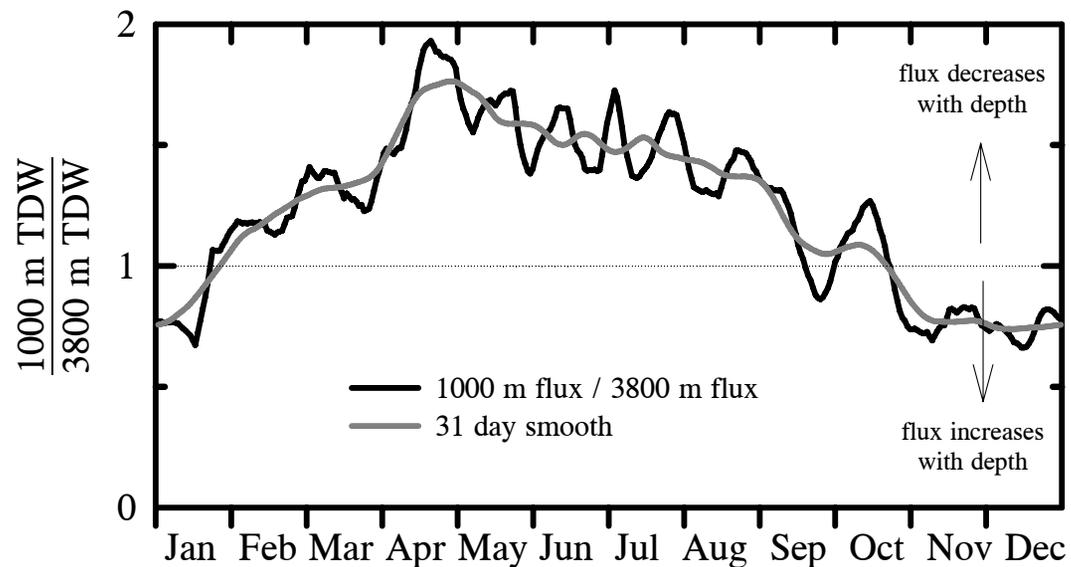
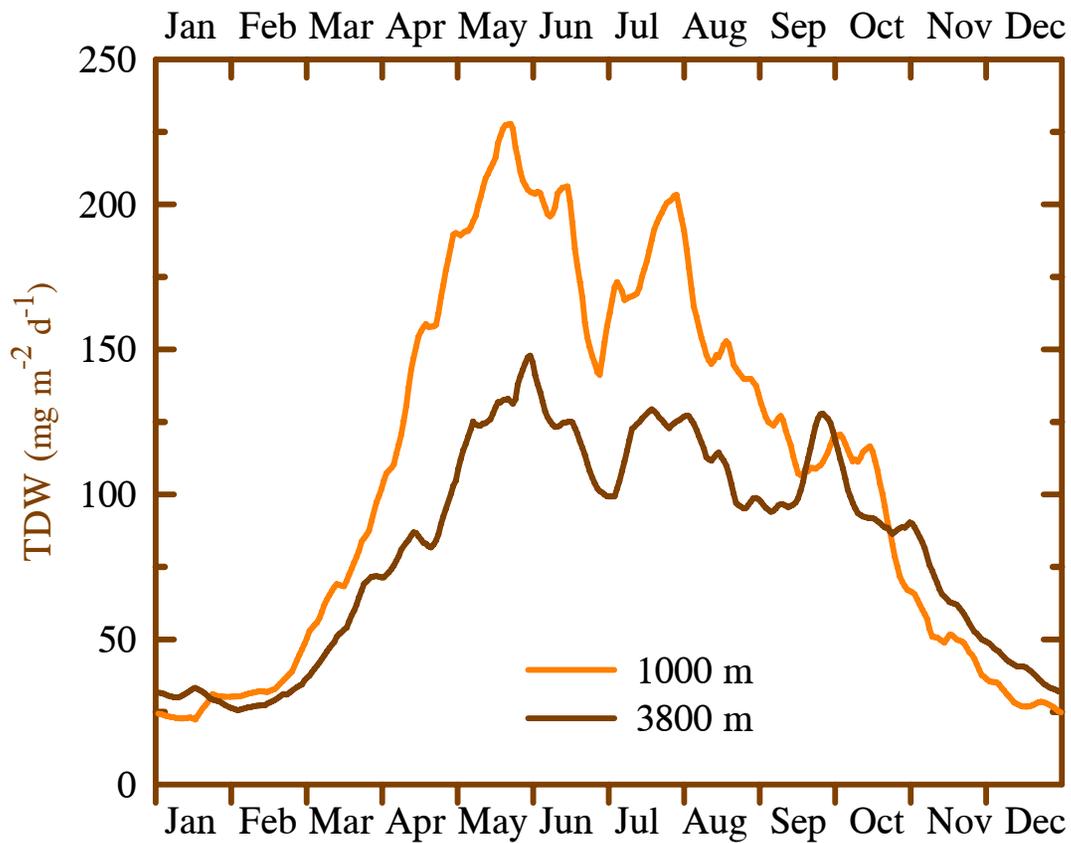
For each constituent, maximum correlation between 1000 m and 3800 m occurred at a 12-16 day shift, which is close to the average deployment period.

Material took ~8 to 24 days to sink 2800 m, for sinking rates of about 100 to 400 m d⁻¹.

The shift for CaCO₃ (12 days) was less than the shift for TDW, BSi and POC (16 days). Is CaCO₃ sinking faster than BSi, and is POC associated with BSi moreso than with CaCO₃?

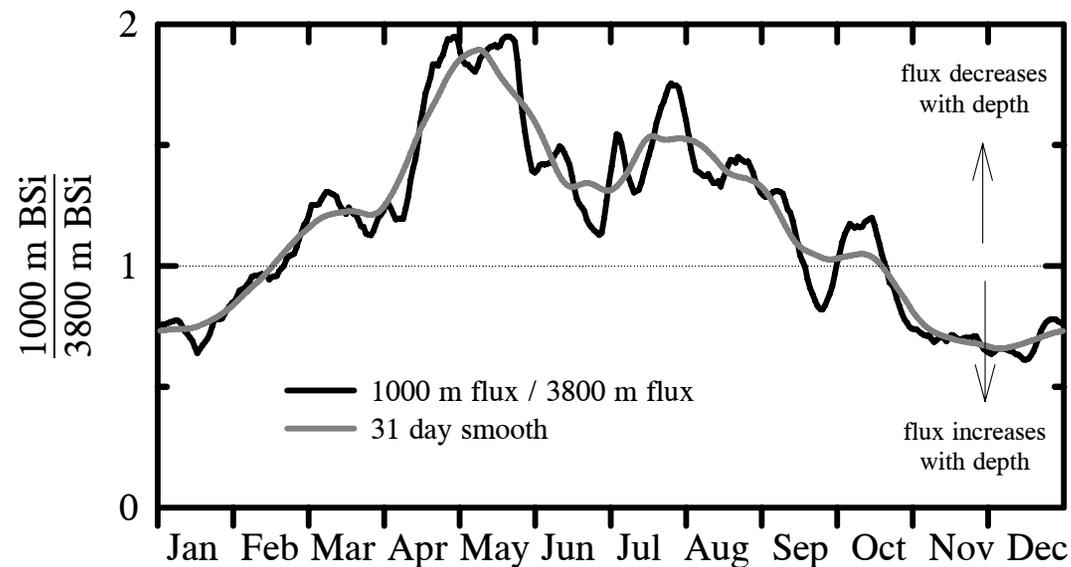
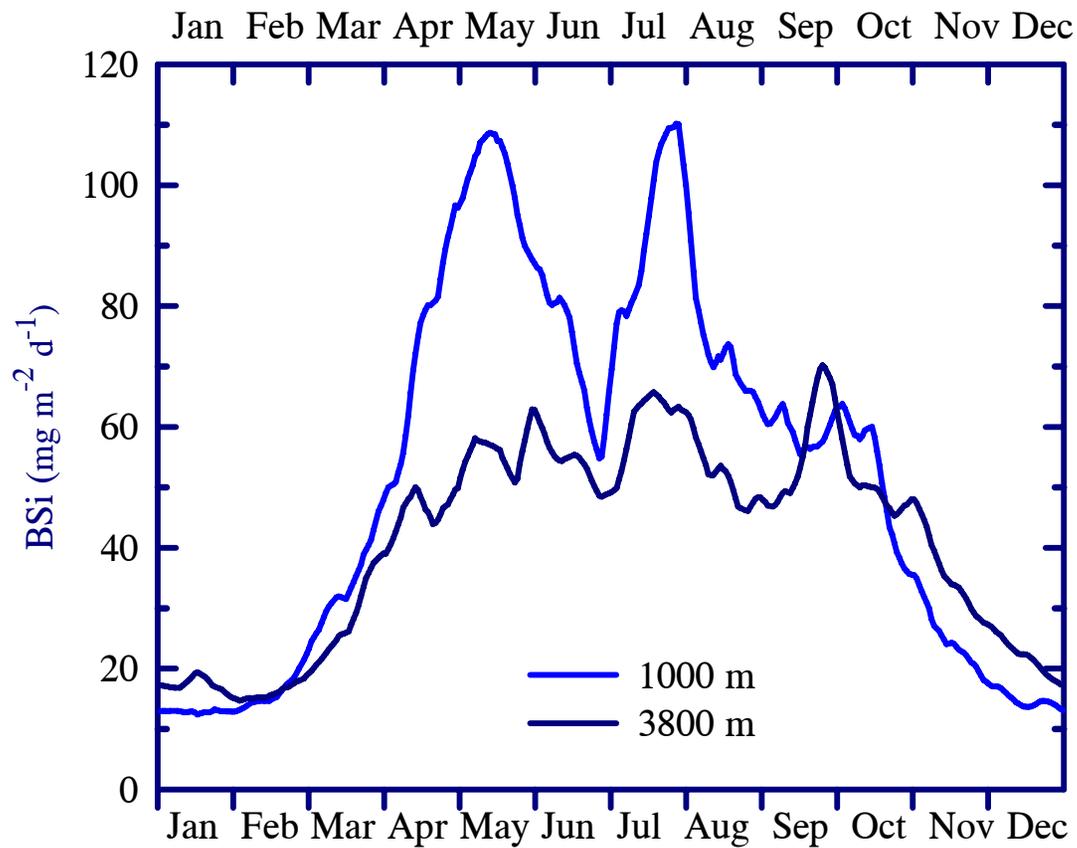
Correlations (lower fig) were determined by shifting the entire time series, not the seasonally averaged fluxes as implied by the upper figure.





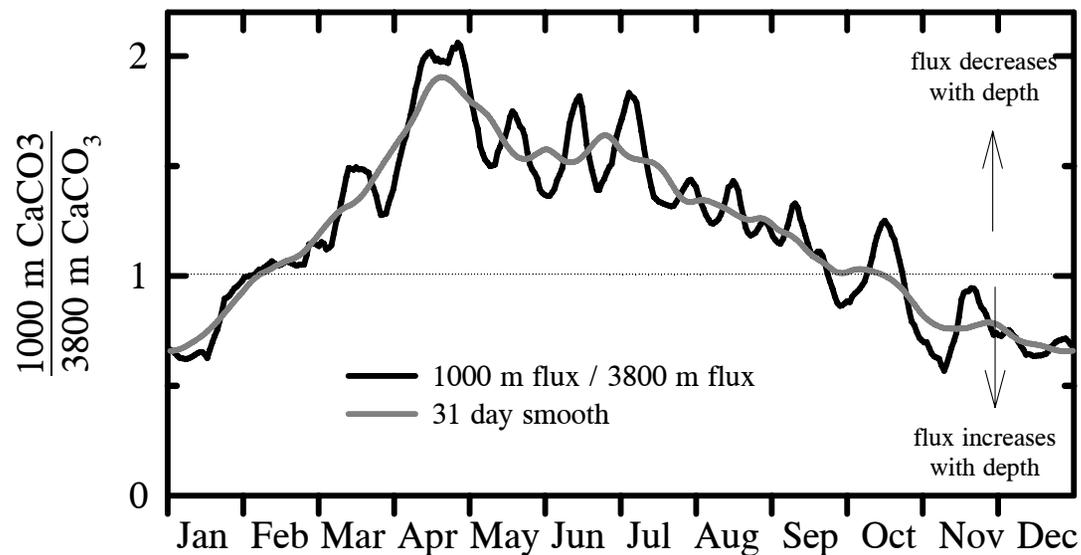
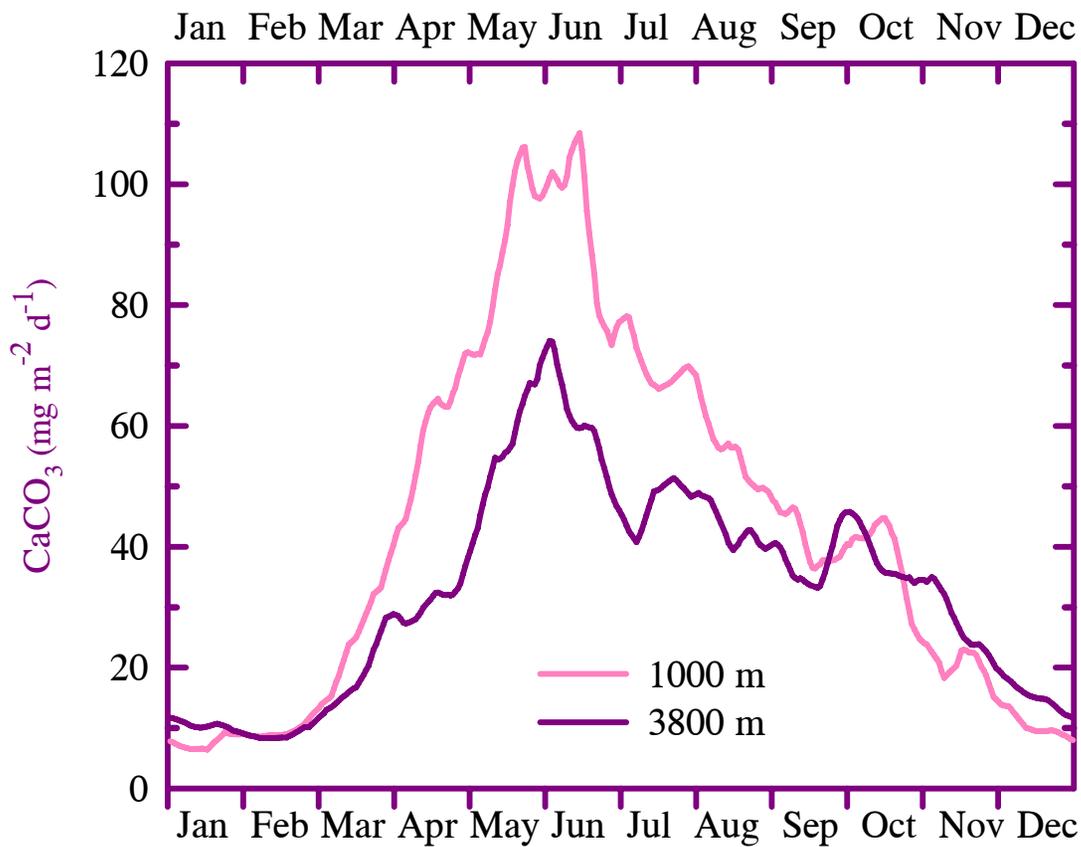
total dry weight flux at OSP
with 16 day shift between
1000 m and 3800 m

seasonally averaged
change in flux
with depth



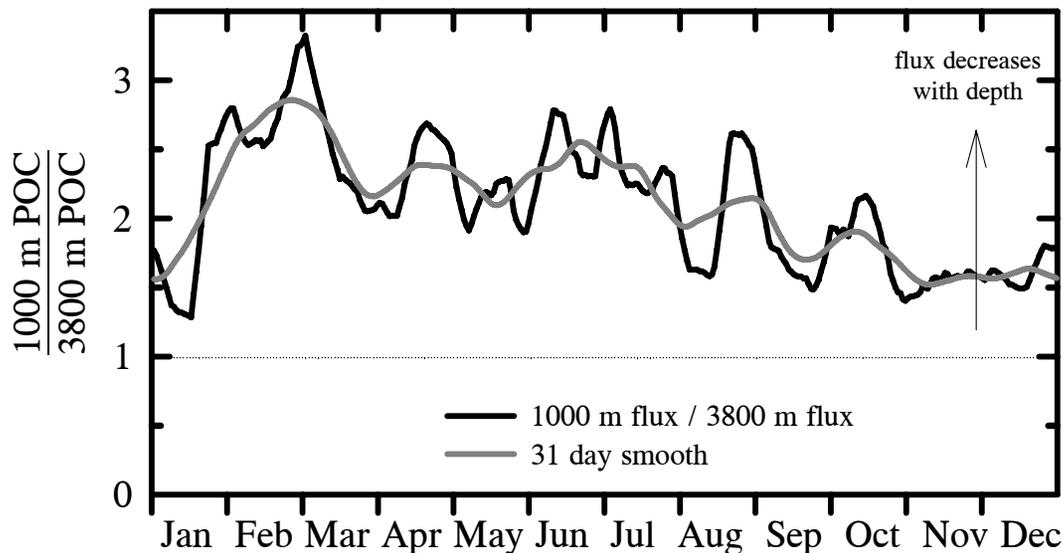
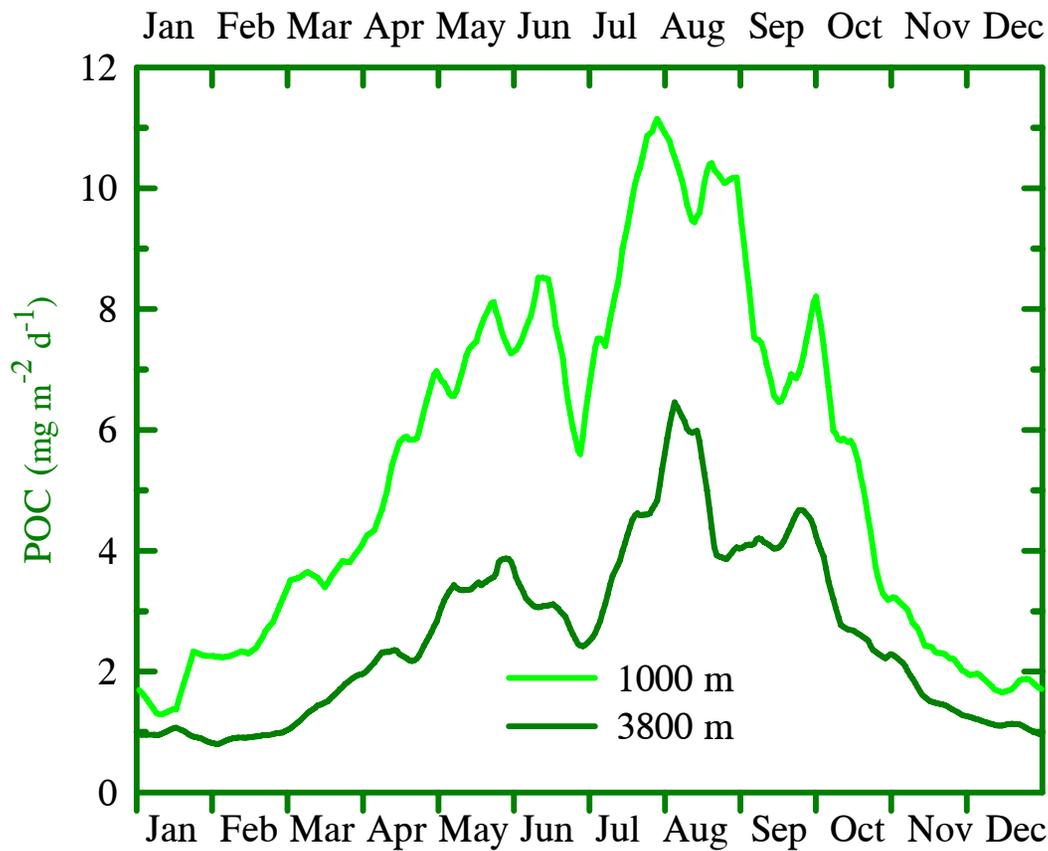
biogenic silica flux at OSP
with 16 day shift between
1000 m and 3800 m

seasonally averaged
change in flux
with depth



CaCO₃ flux at OSP
with 12 day shift between
1000 m and 3800 m

seasonally averaged
change in flux
with depth



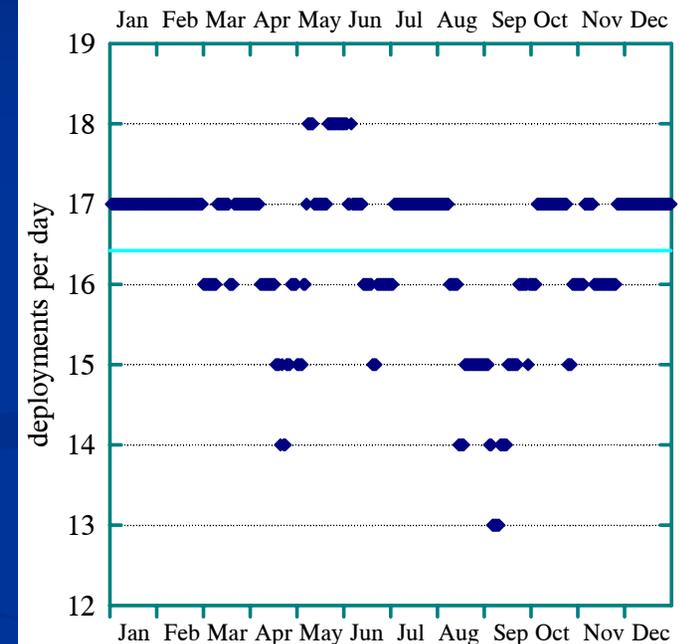
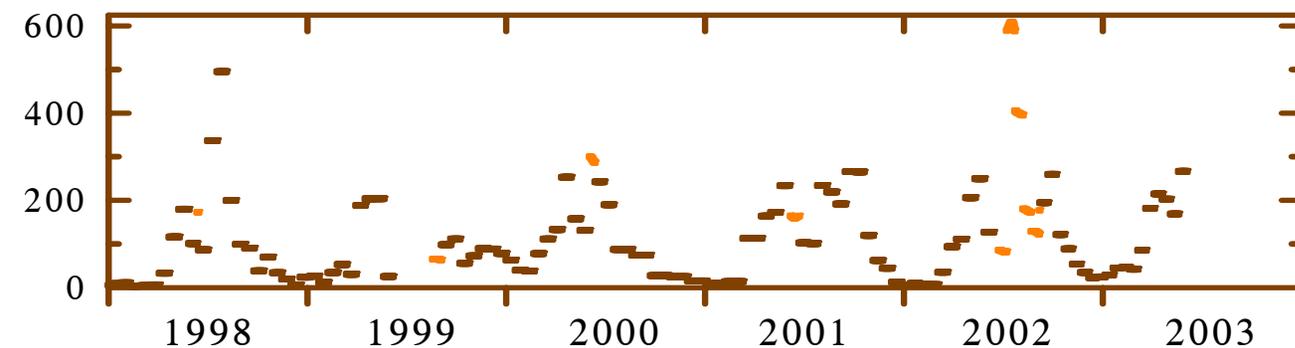
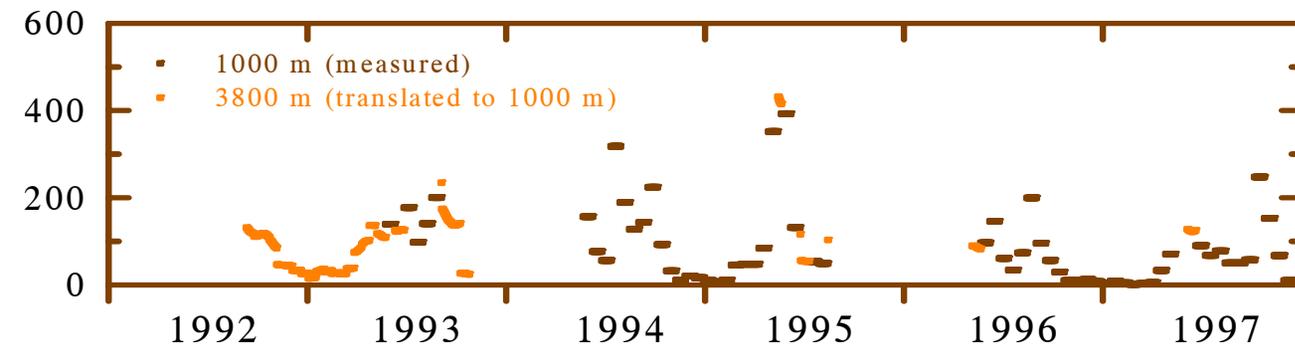
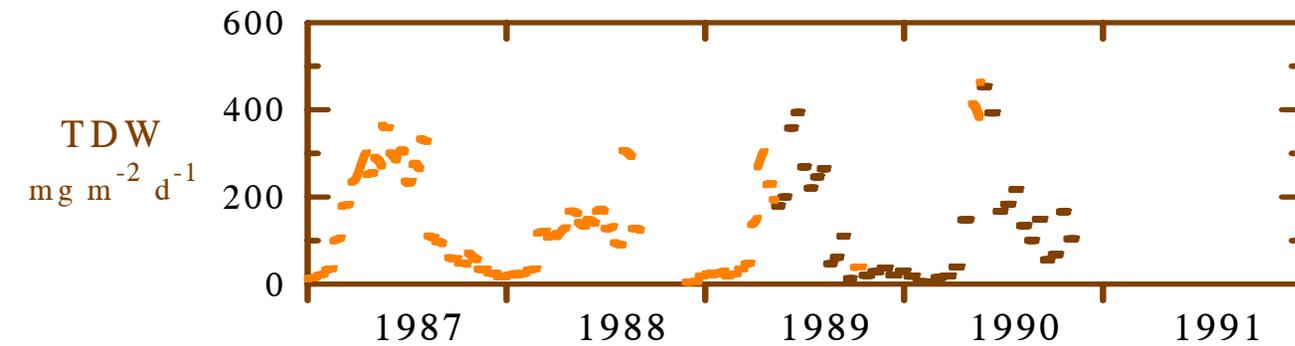
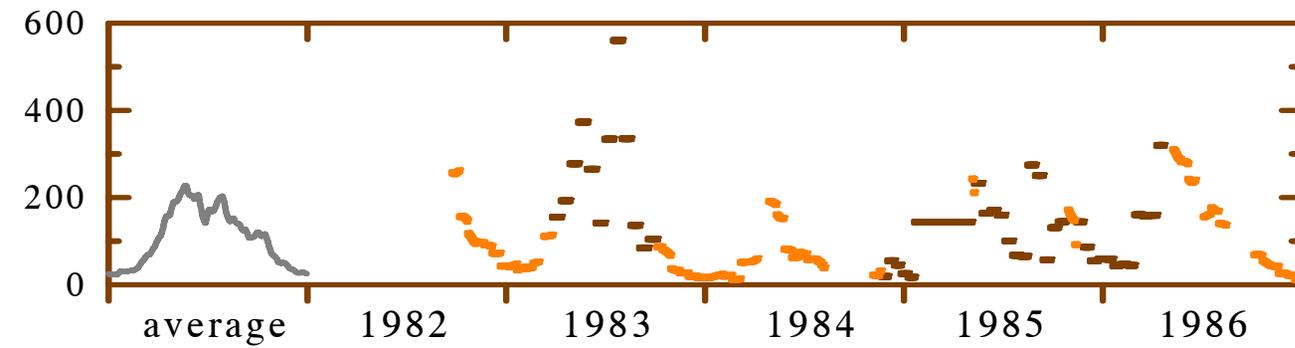
POC flux at OSP with 16 day shift between 1000 m and 3800 m

Shoaling of seasonal thermocline
in Aug \rightarrow high primary production
(D Crawford, pers. comm.)

similar double peak but different
remineralsation dynamics than
presented yesterday at HOTS

seasonally averaged
change in flux
with depth

1000 m fluxes at OSP with merged 3800 m fluxes



16+ years

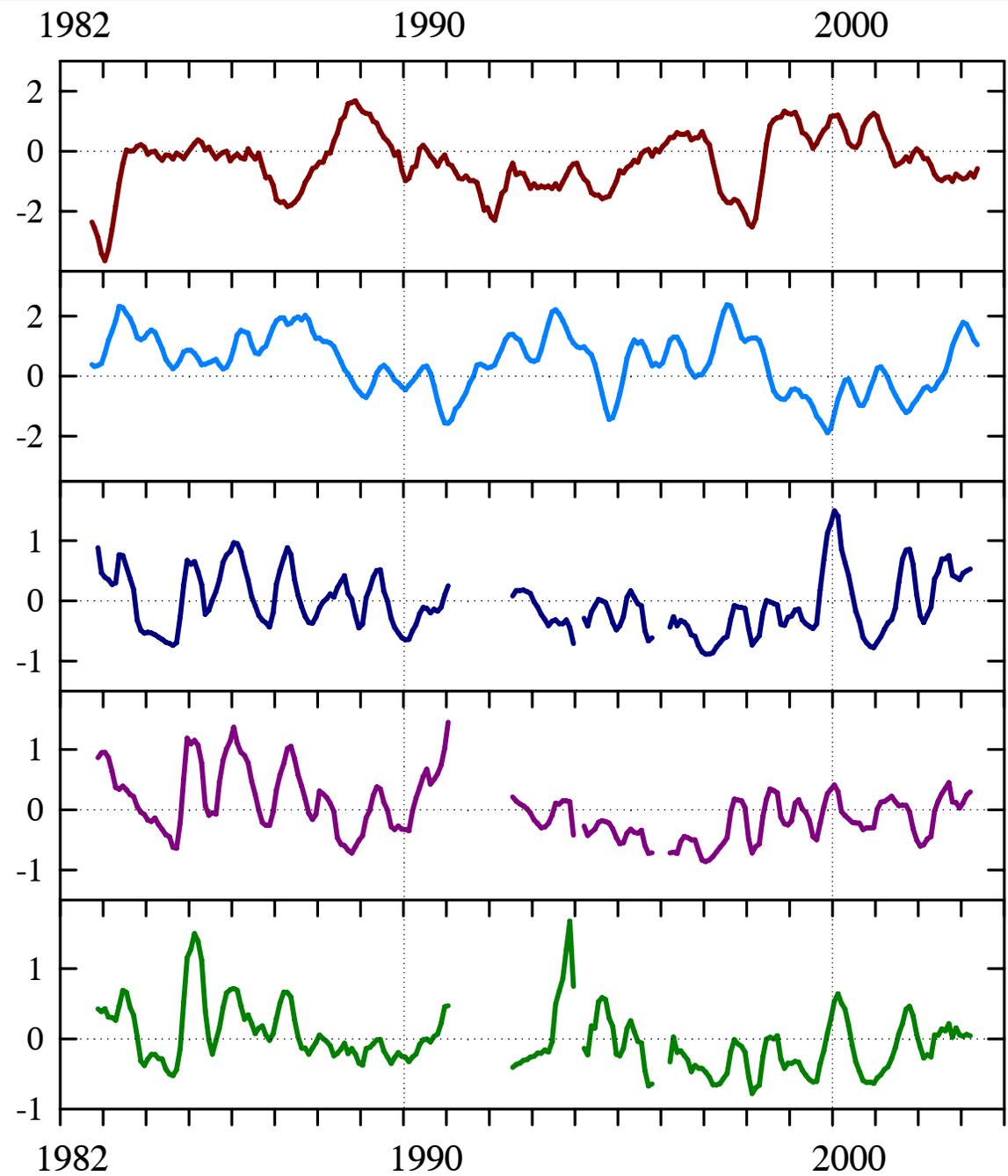
Southern Oscillation
Index

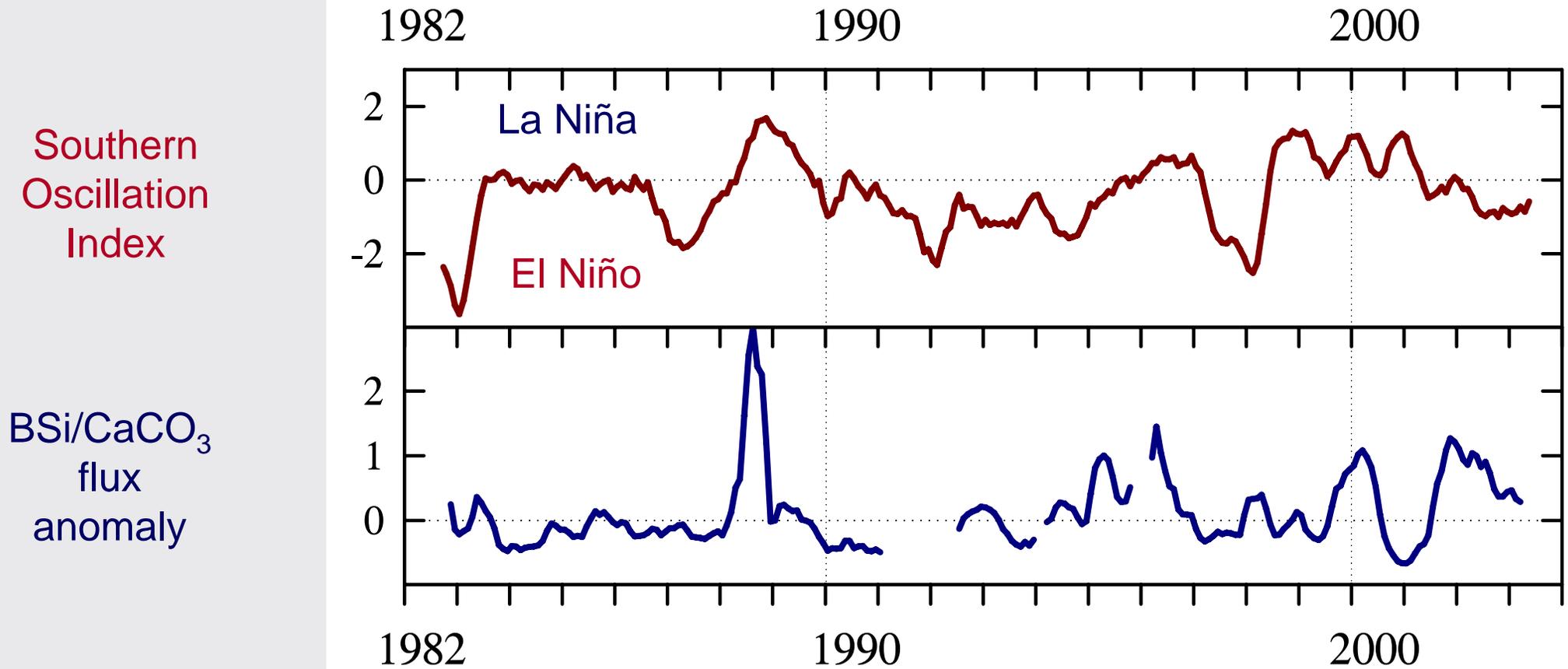
Pacific Decadal
Oscillation

BSi flux
anomaly

CaCO₃ flux
anomaly

POC flux
anomaly

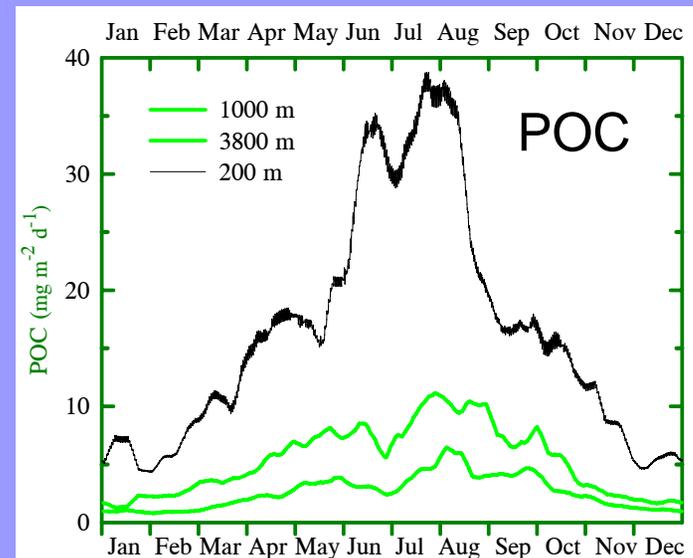
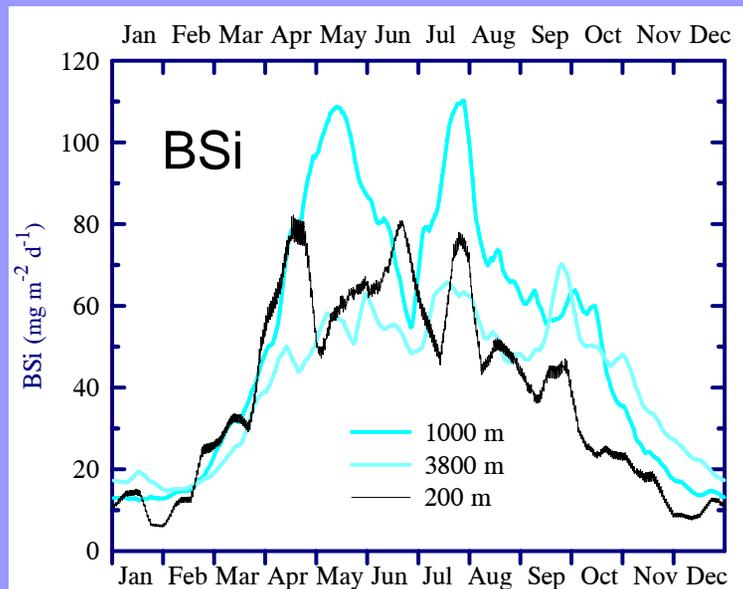
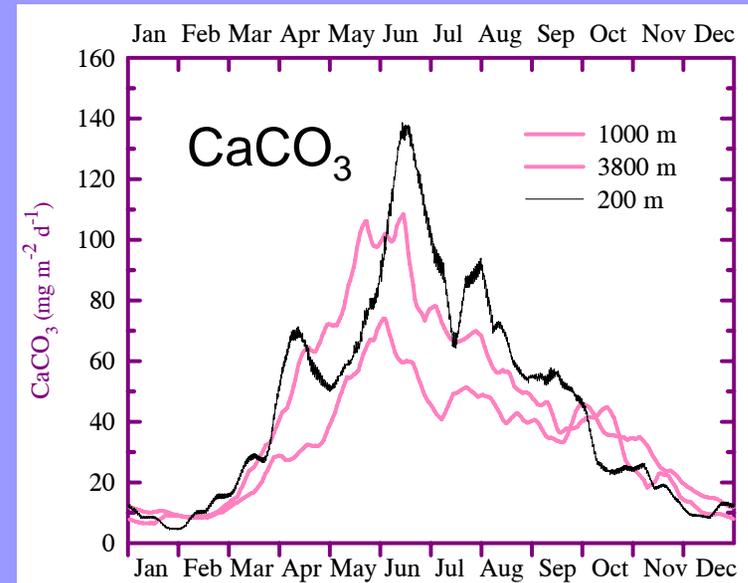
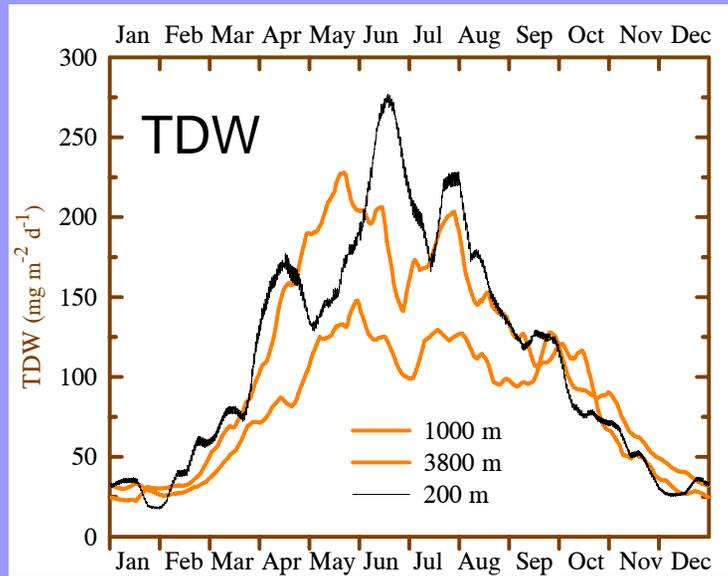




CONCERNS and ONGOING WORK

1. Large anomalies in general. Need to average fluxes over longer periods?
How long?
2. Large anomalies in winter. Longer averaging?
3. 1988 is a period when 3800 m fluxes were translated to 1000 m.
Need to assure 1988 anomaly exists relative to the 3800 m time-series.

200 m fluxes at OSP



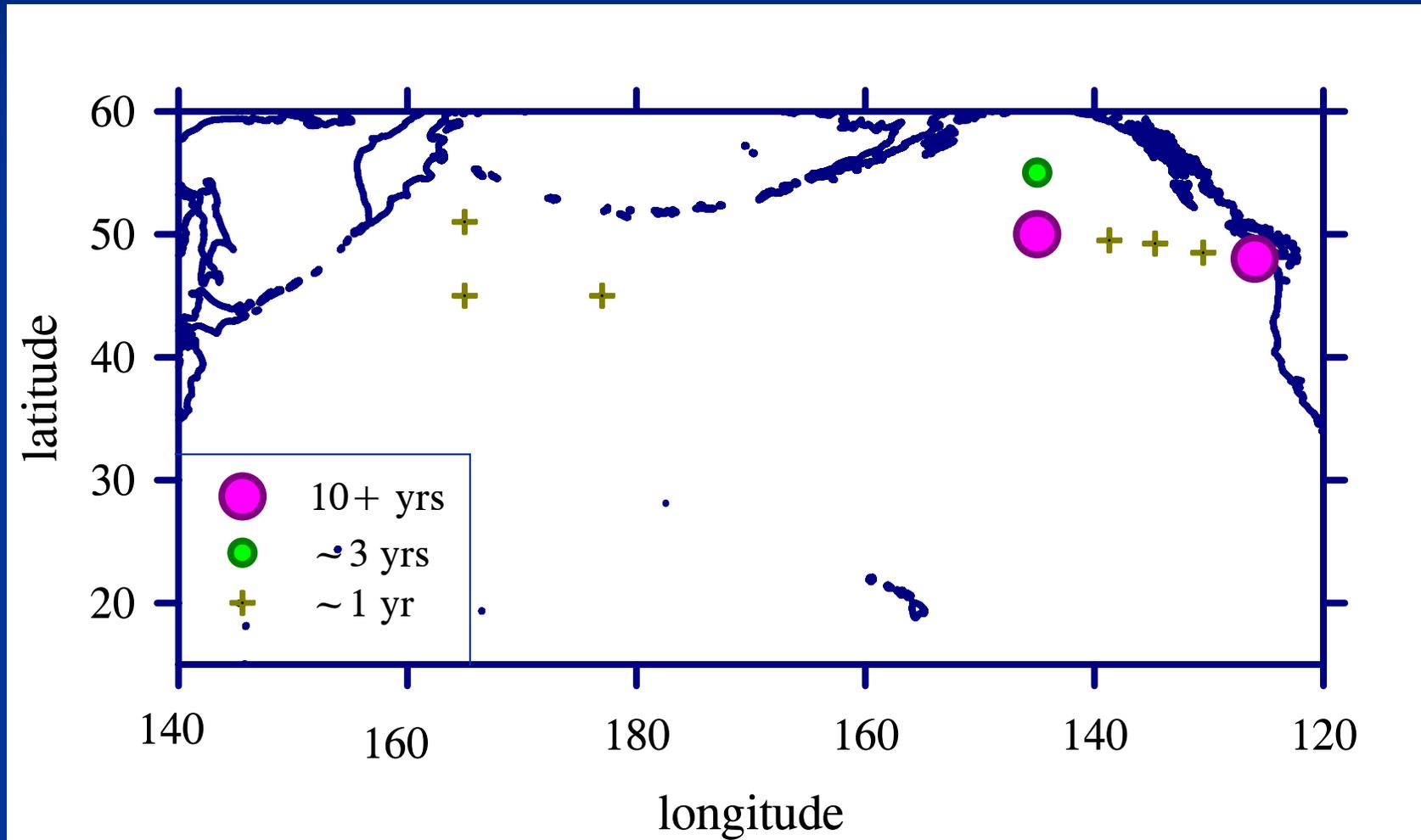
OSP conclusions

- The seasonal cycle of flux at OSP occurs with spring and summer peaks in the fluxes of BSi and POC, but only a spring peak in the flux of CaCO_3 .
- Particles sink at rates of $100\text{-}400 \text{ m d}^{-1}$ below 1000 m, with the possibility CaCO_3 sinks faster than BSi and POC.
- POC follows BSi more closely than it follows CaCO_3 .
- The time series from OSP is long enough for correlations with SOI and PDO. Preliminary analysis shows a greater predominance of diatoms during La Niña and coccolithophorids during El Niño, consistent with Wong and Crawford (2002).

TIME SERIES FROM:

- Ocean Station PAPA
- Alaska Gyre and North Pacific
Western Gyre
- Line P with emphasis at La Perouse Bank

Alaska Gyre: 1990 to 1992



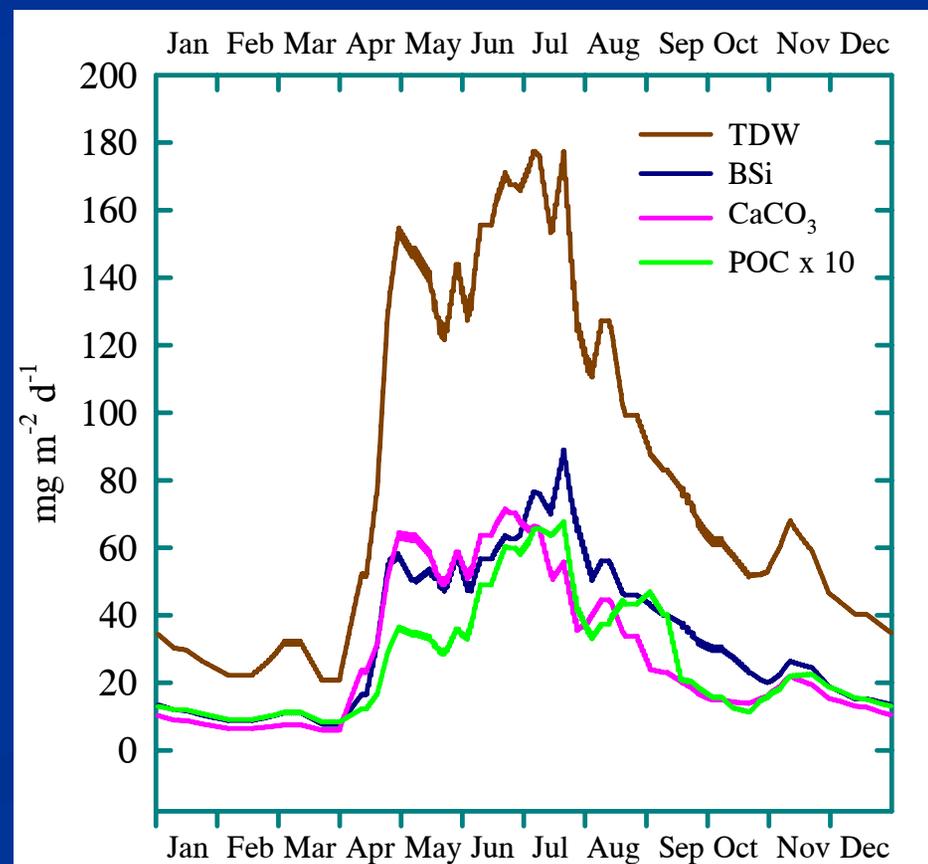
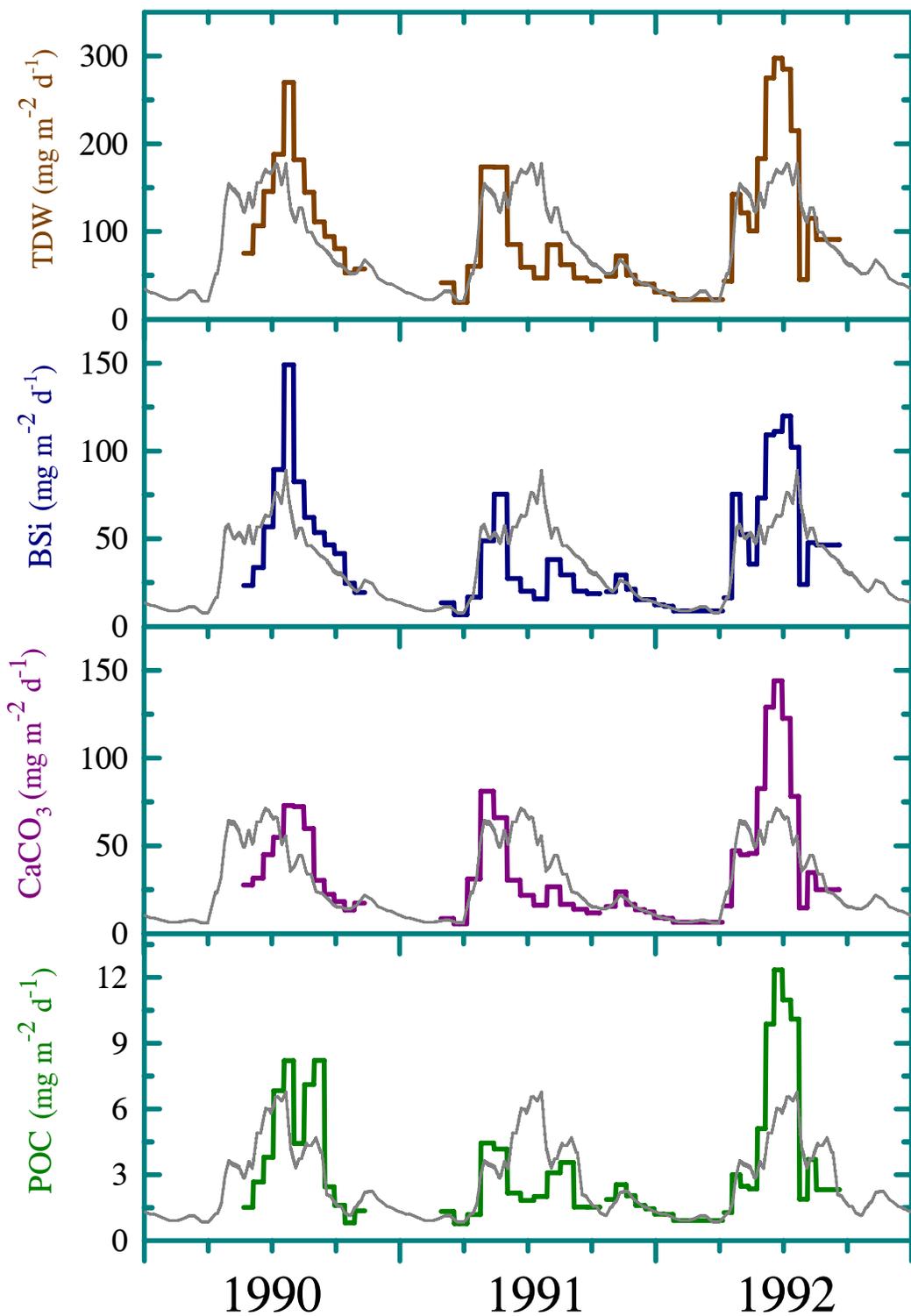
Alaska Gyre

55° 00' N

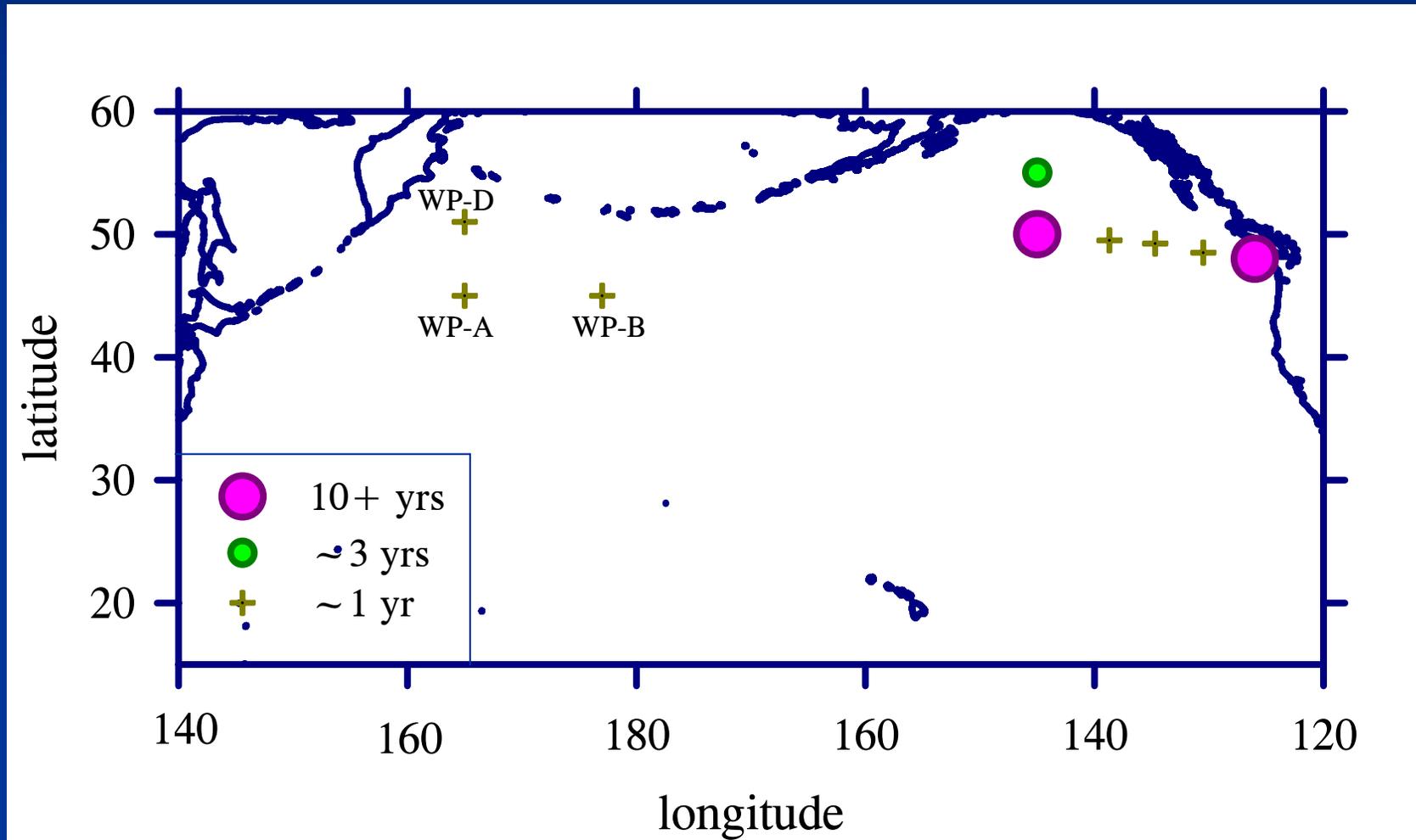
145° 00' W

trap depth = 3700 m

bottom depth = 3900 m



3 stations in the Western Pacific August, 1991 to June, 1992



stations WP aka GEMS

North Pacific Western Gyre sediment-trap fluxes

WP D

51° 25' N
165° 13' E
4300 m

bottom depth = 5200 m

WP A

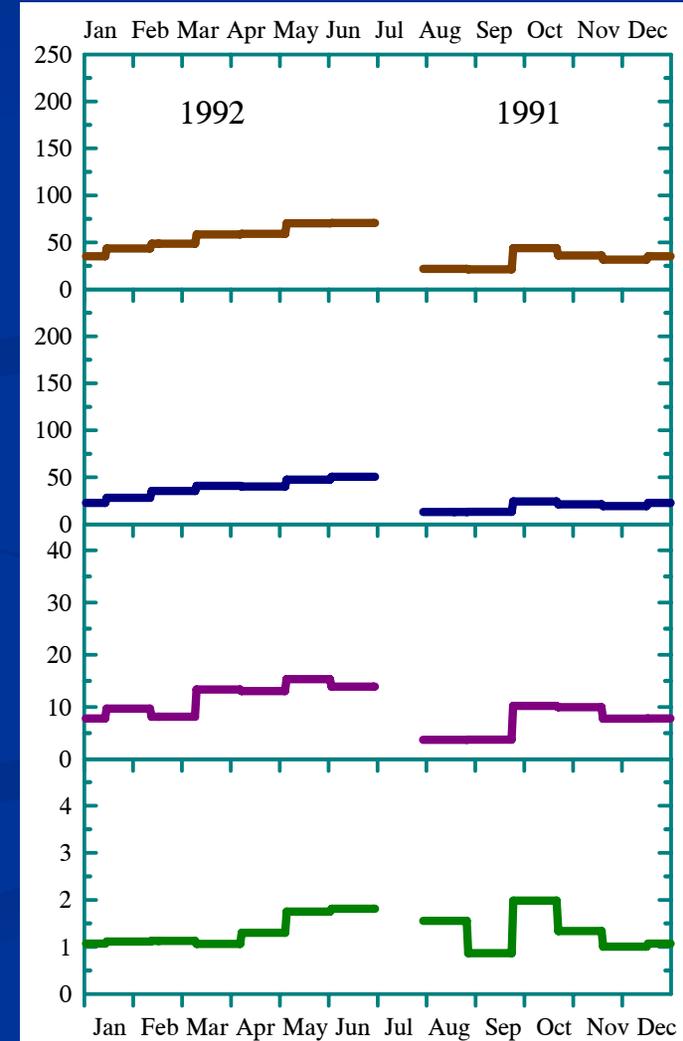
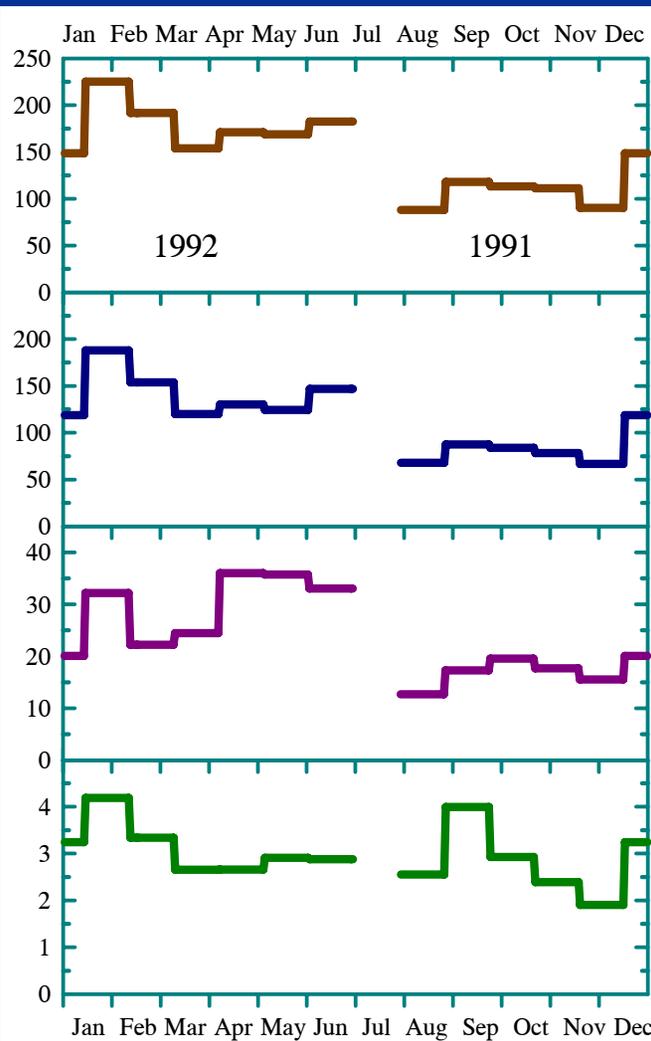
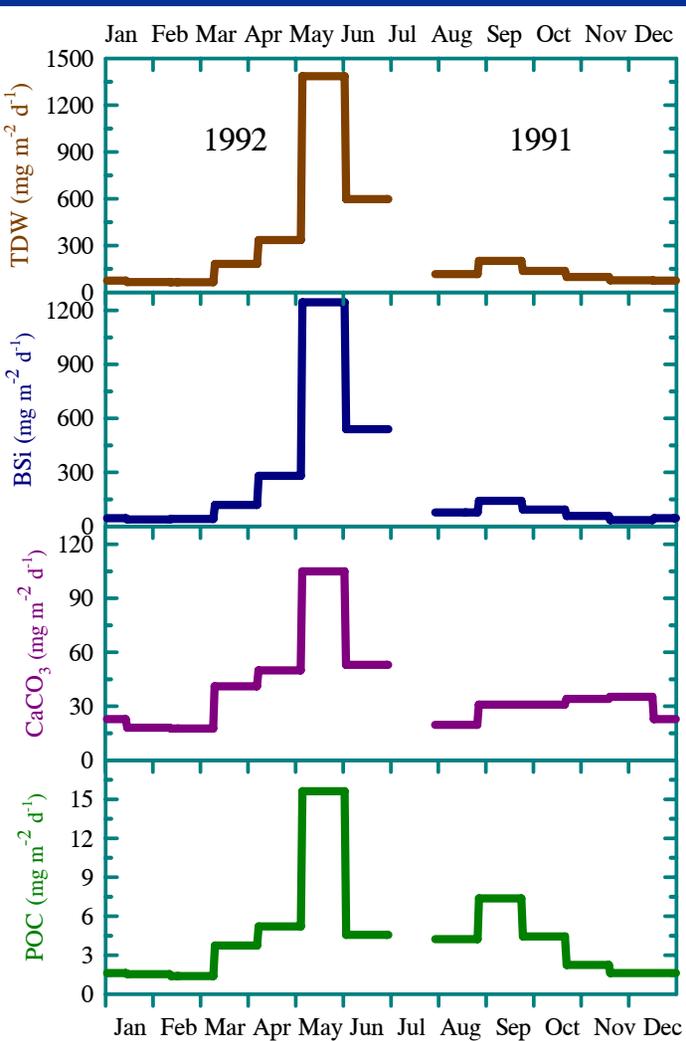
44° 58' N
165° 00' E
5300 m

bottom depth = 5860 m

WP B

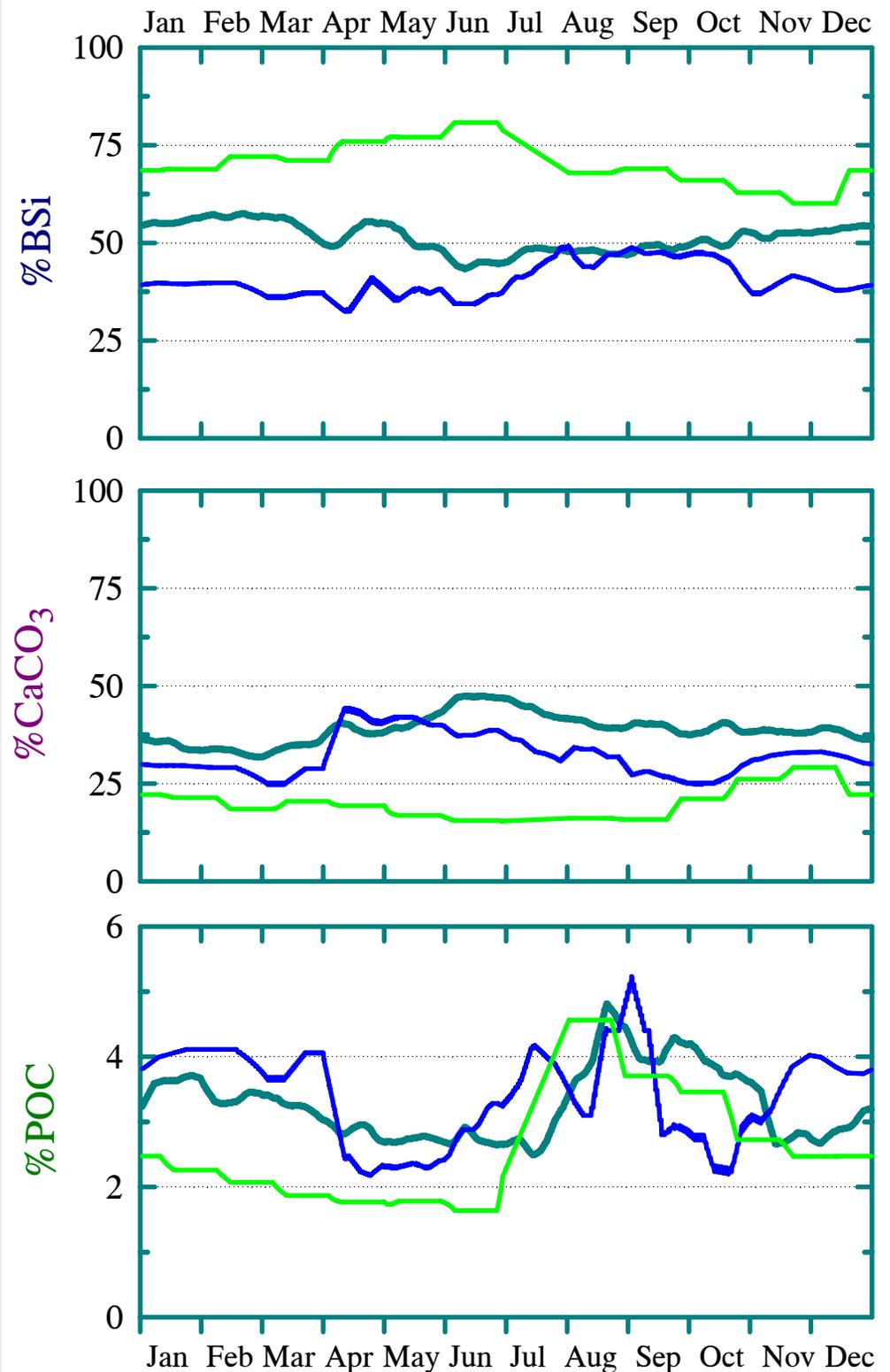
45° 05' N
176° 53' W
5200 m

bottom depth = 5755 m



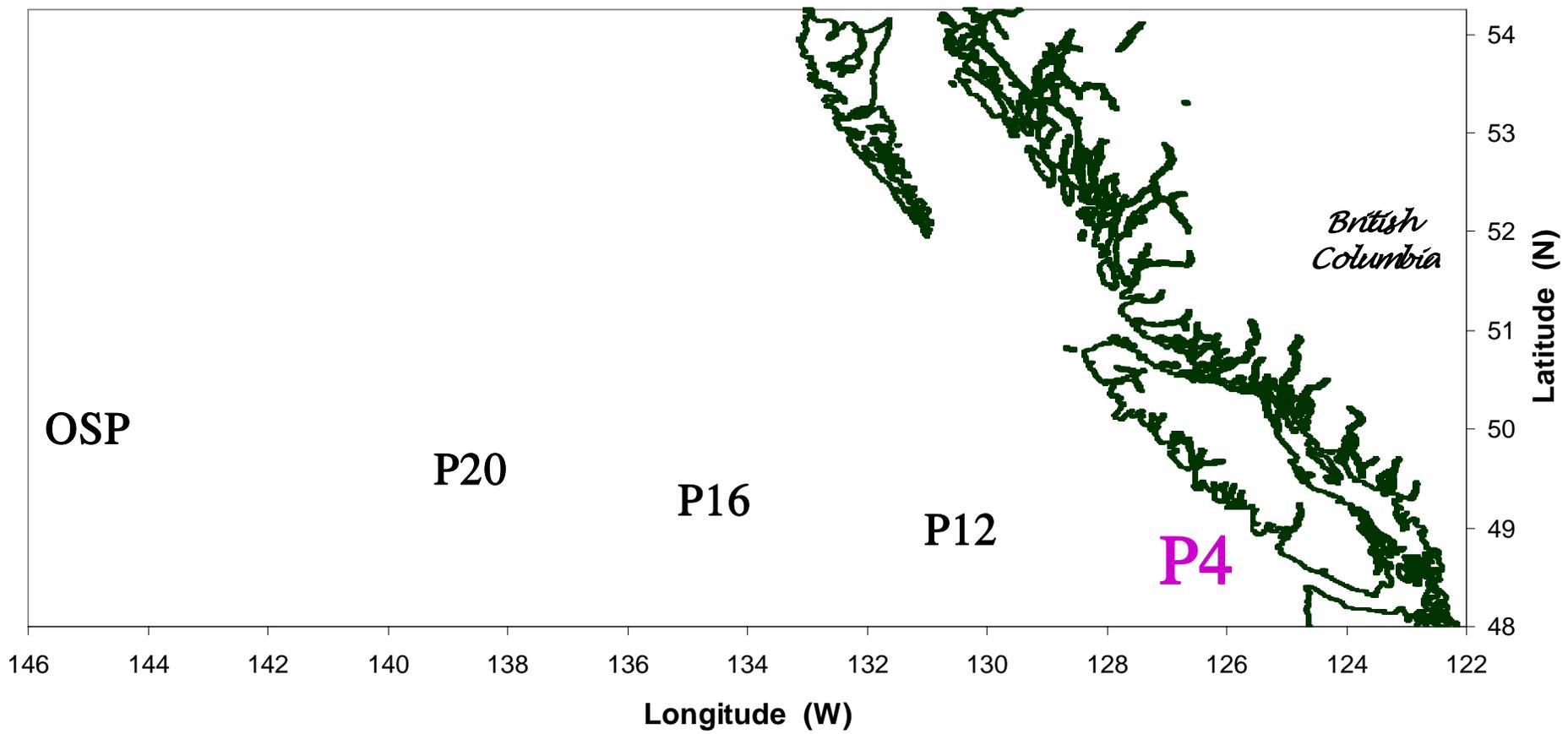
sediment composition at deep traps

- W Pacific
4300 m – 5300 m
- OSP
3800 m
- Alaska Gyre
3700 m



TIME SERIES FROM:

- Ocean Station PAPA
- Alaska Gyre and North Pacific Western Gyre
- Line P with emphasis at
La Perouse Bank



La Perouse Bank

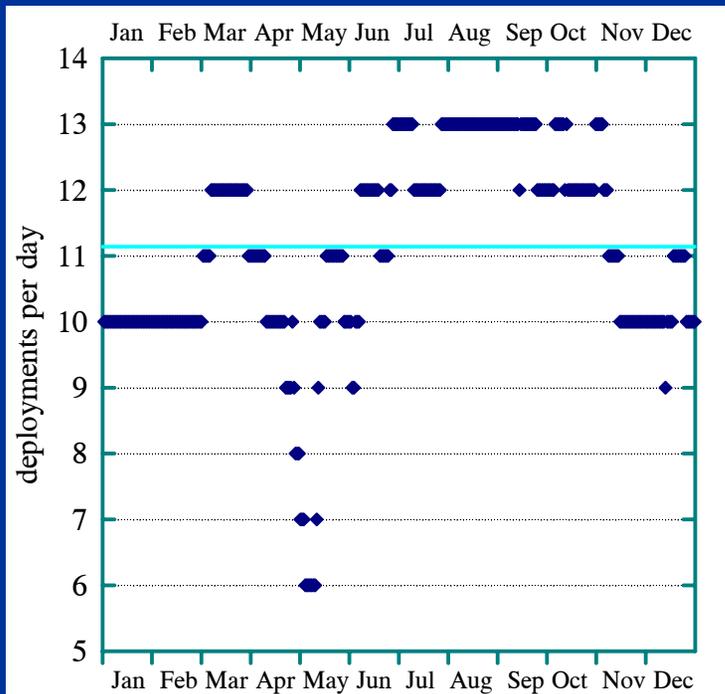
48° 35' N

126° 24' W

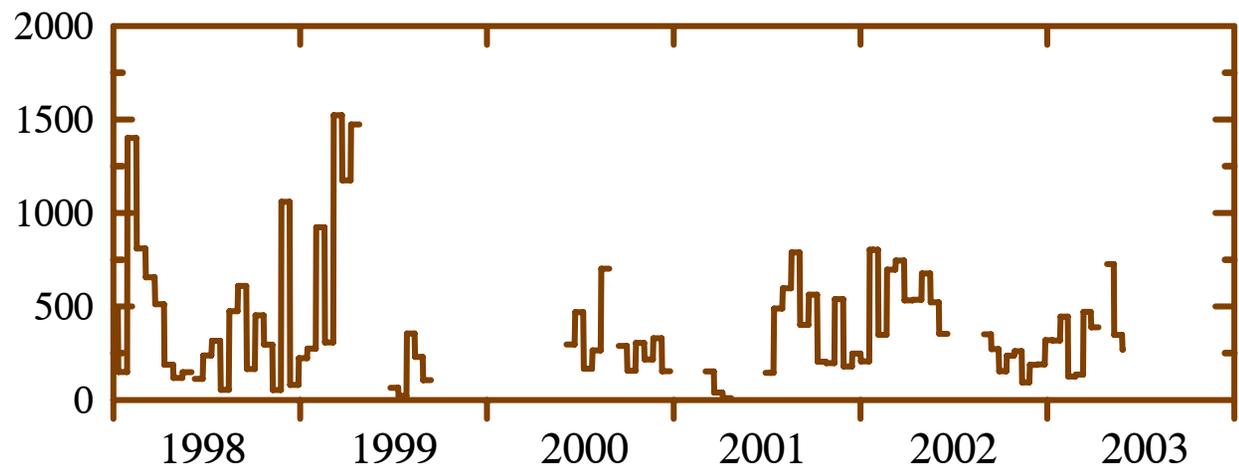
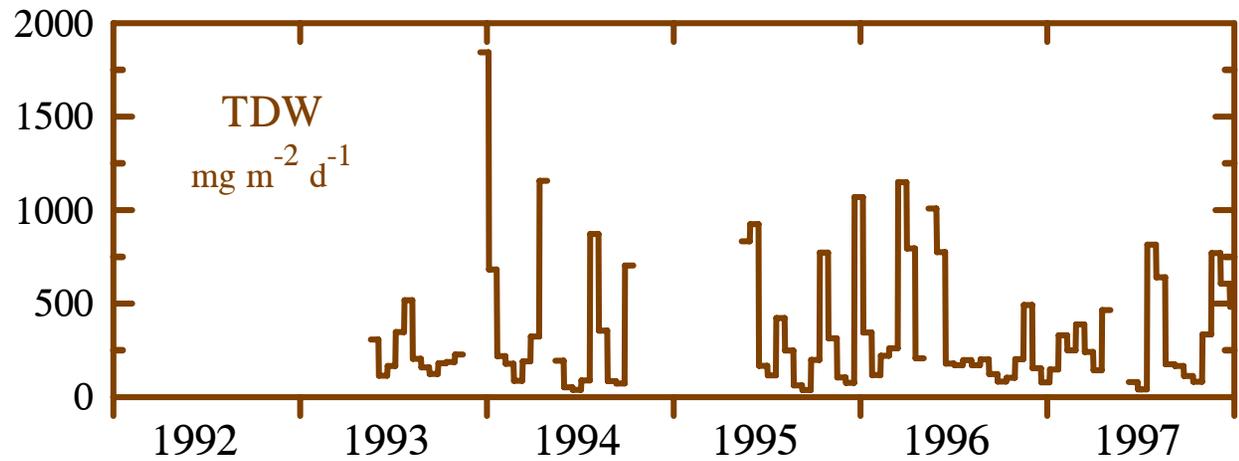
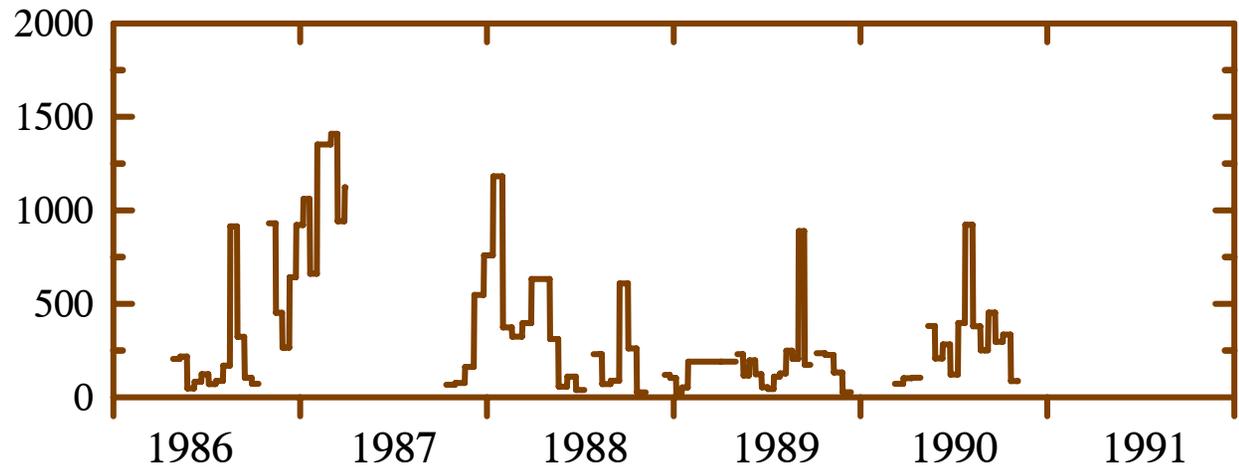
prior to 1991: 500 m

after 1991: 700 m

bottom depth = 1300 m



11 yrs of data



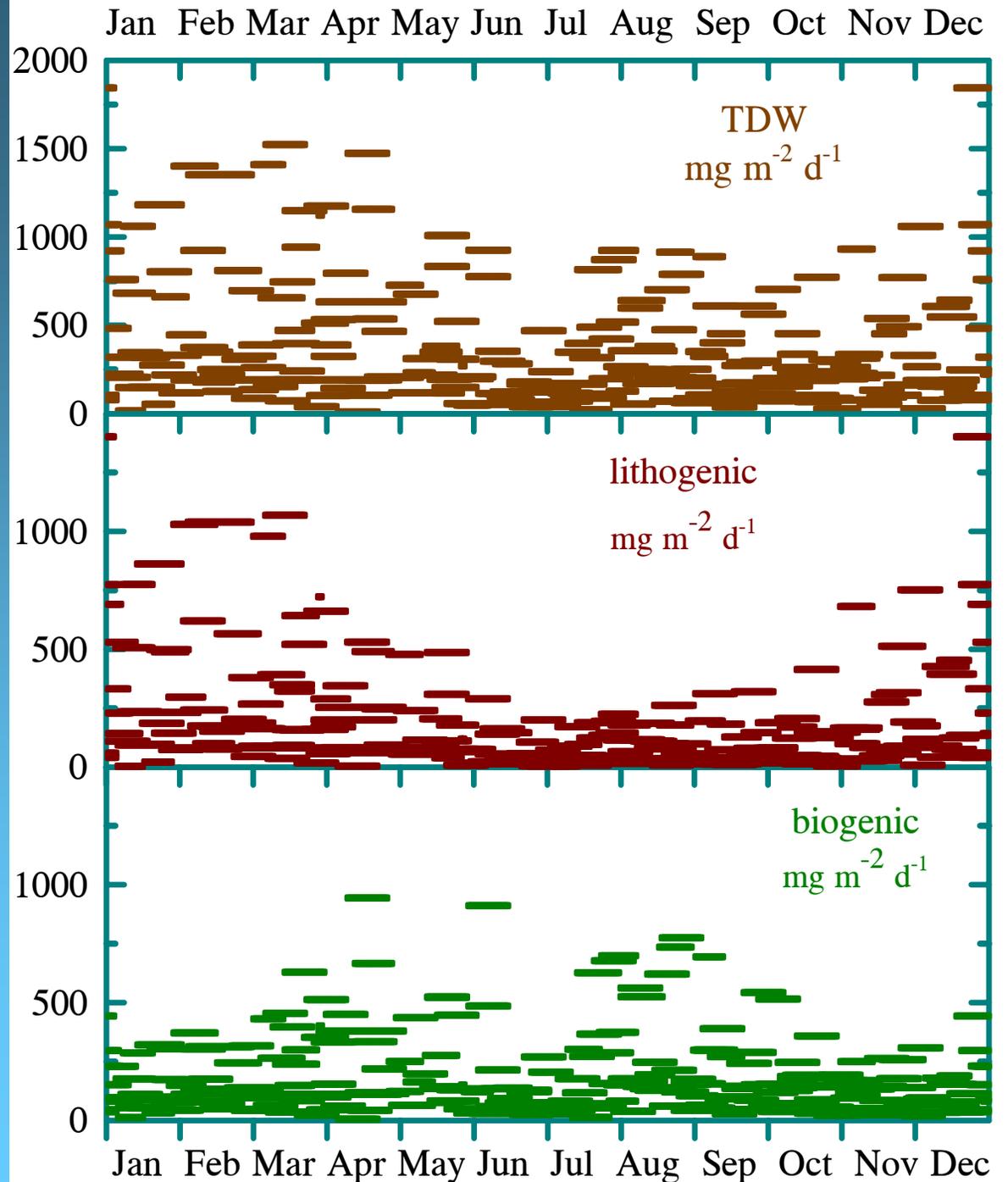
Resuspended fluxes at La Perouse Bank



no signal of primary production in the bulk flux at La Perouse Bank, and up to 75% of the sediment was lithogenic.

lithogenic sediment = total - biogenic

biogenic sediment = $\text{BSi} + \text{CaCO}_3 + 1.85 (\text{POC})$



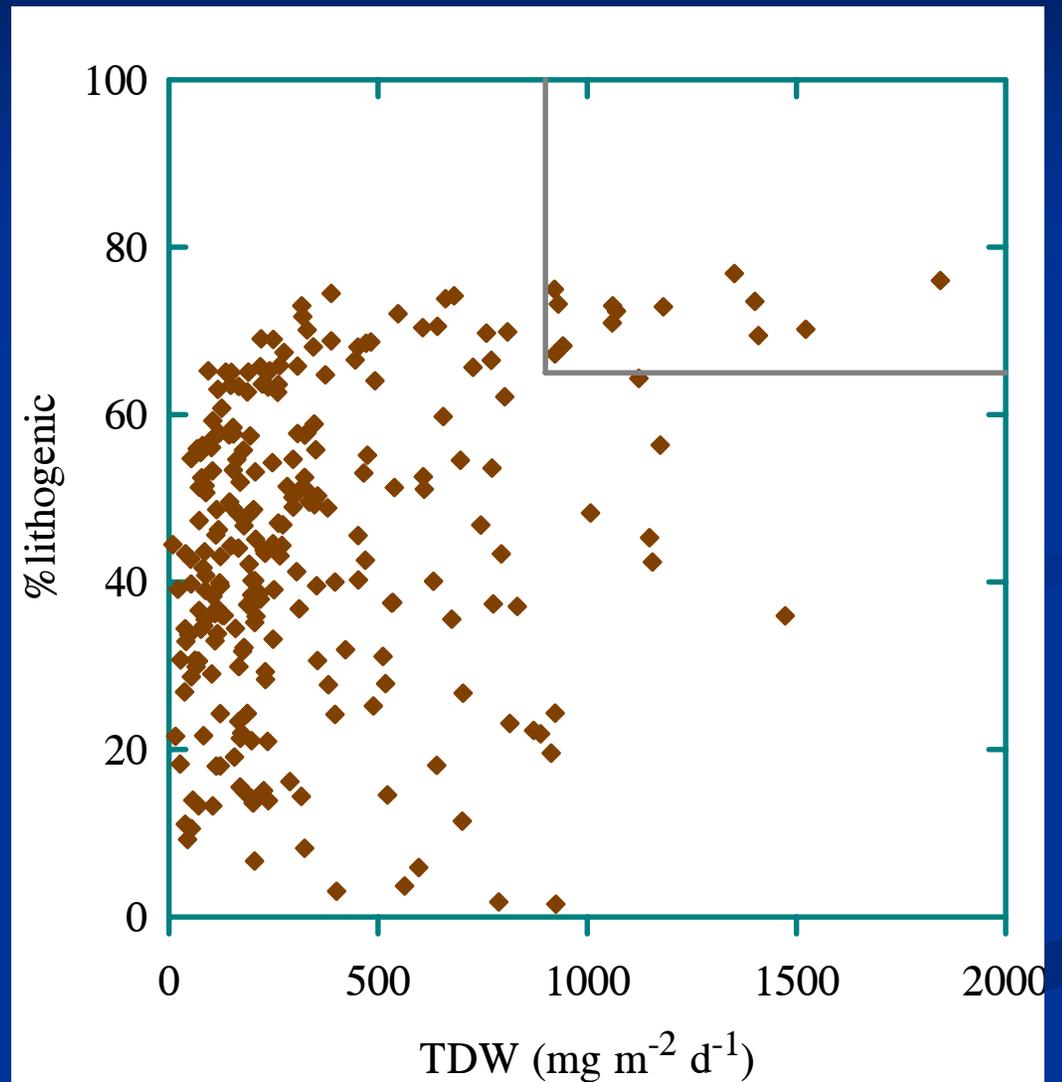
Resuspended sediment at La Perouse Bank

The composition of resuspended sediment was modelled after the samples where fluxes were more than 65% lithogenic and the total flux was $> 900 \text{ mg m}^{-2} \text{ d}^{-1}$

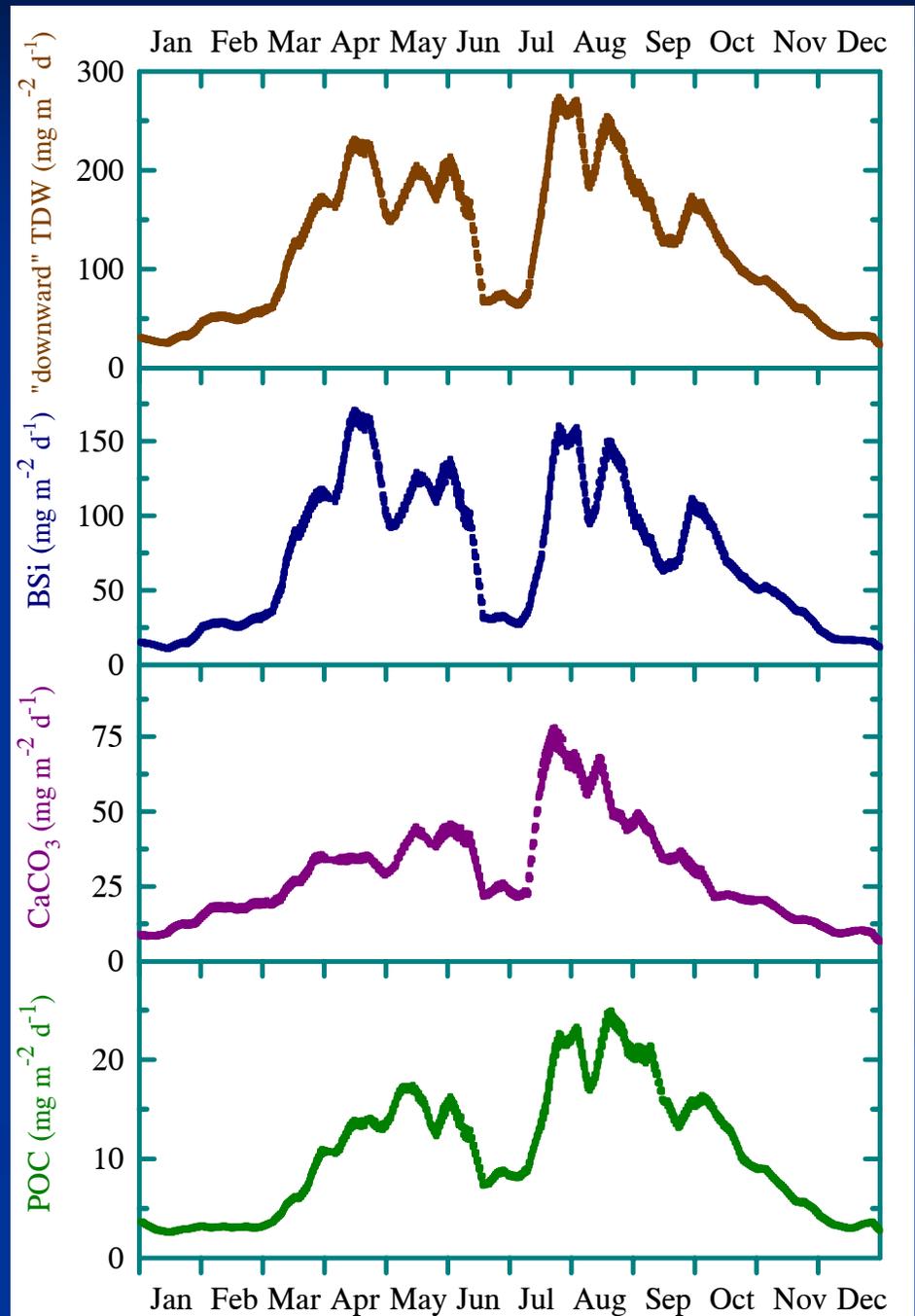
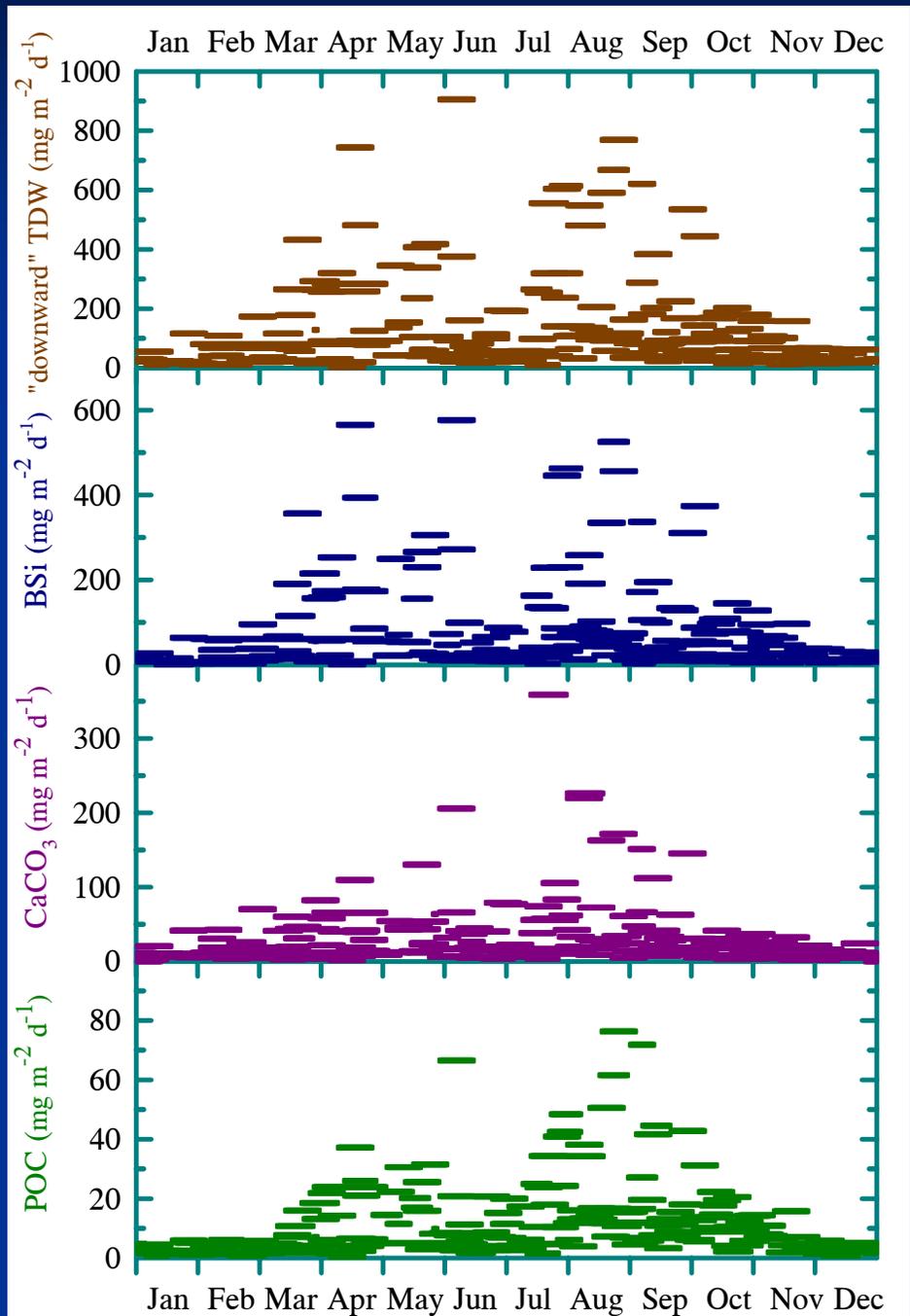
the mean composition of these samples was:

73% lithogenic
13% biogenic silica
8.3% CaCO_3
3.5% POC

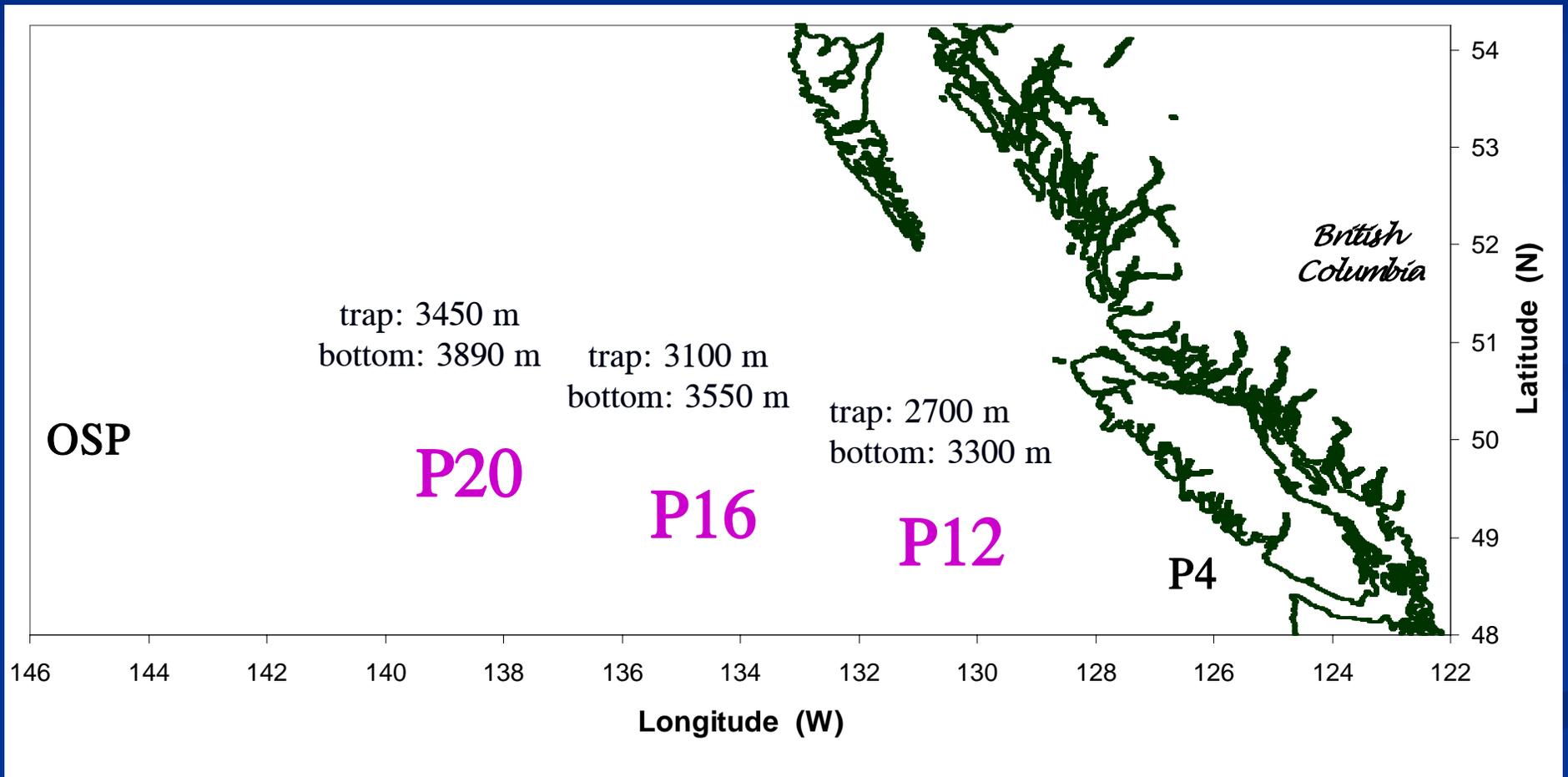
These were all in the winter when the “downward” flux would have been minimal.



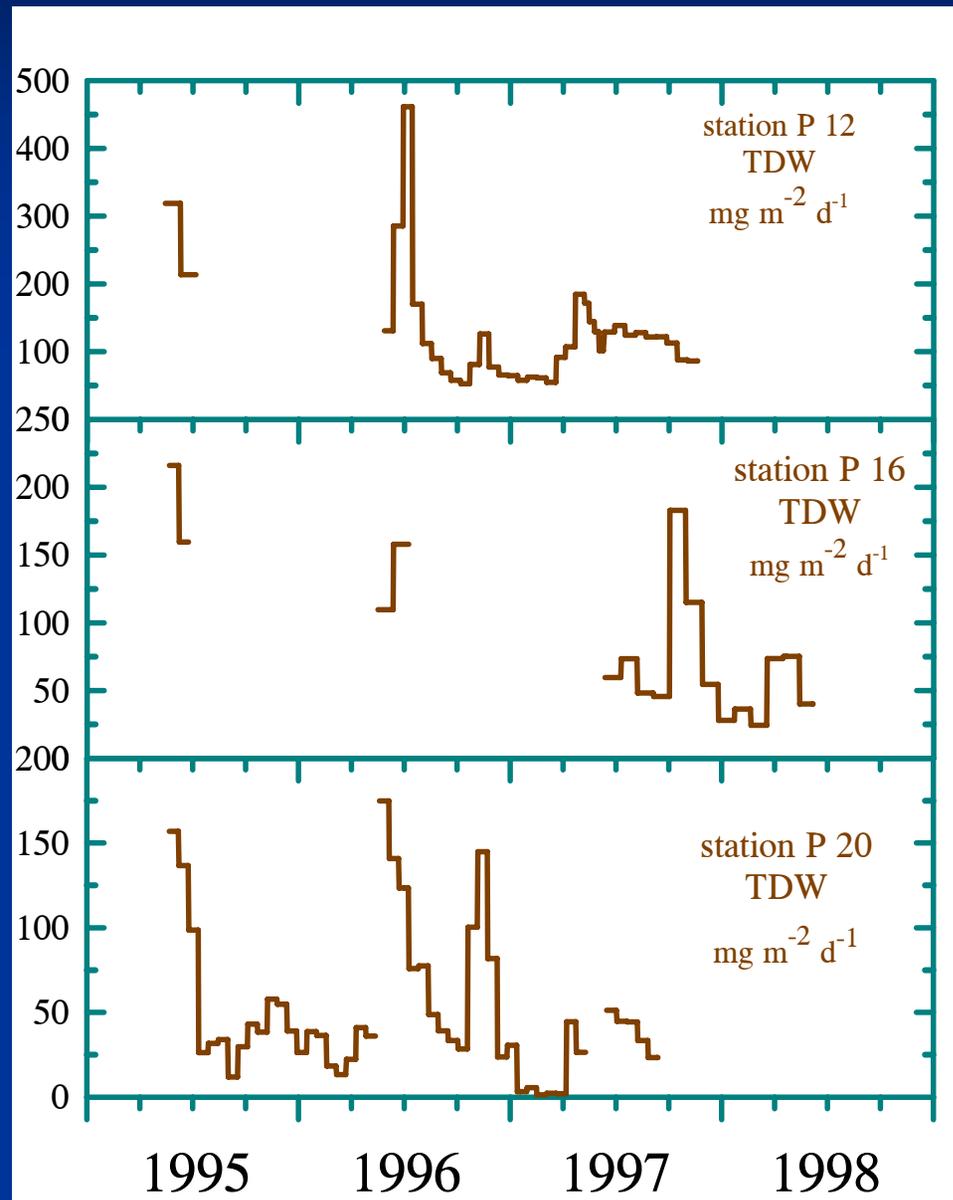
The downward flux (total – resuspended) at La Perouse Bank



compilation from Line P

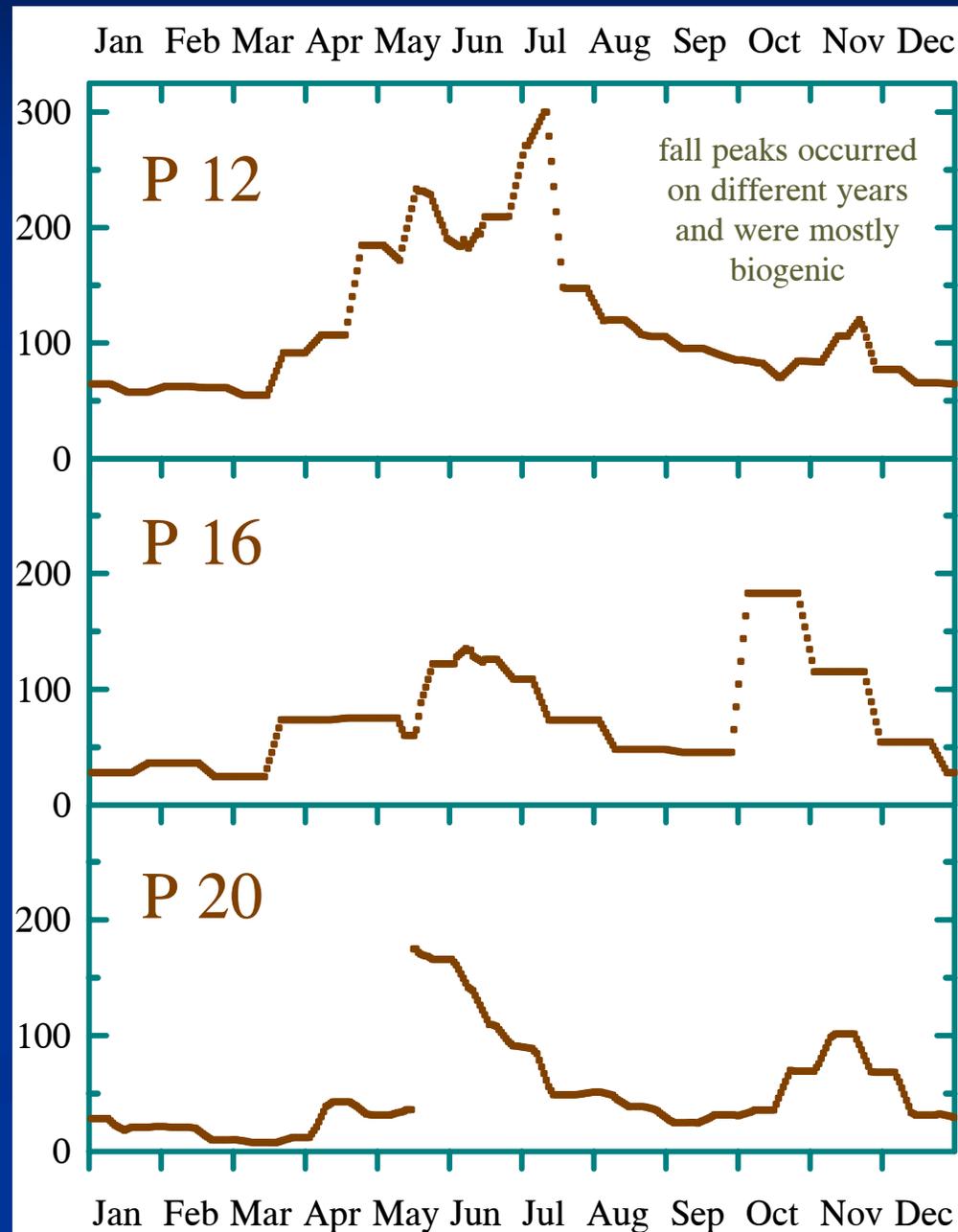
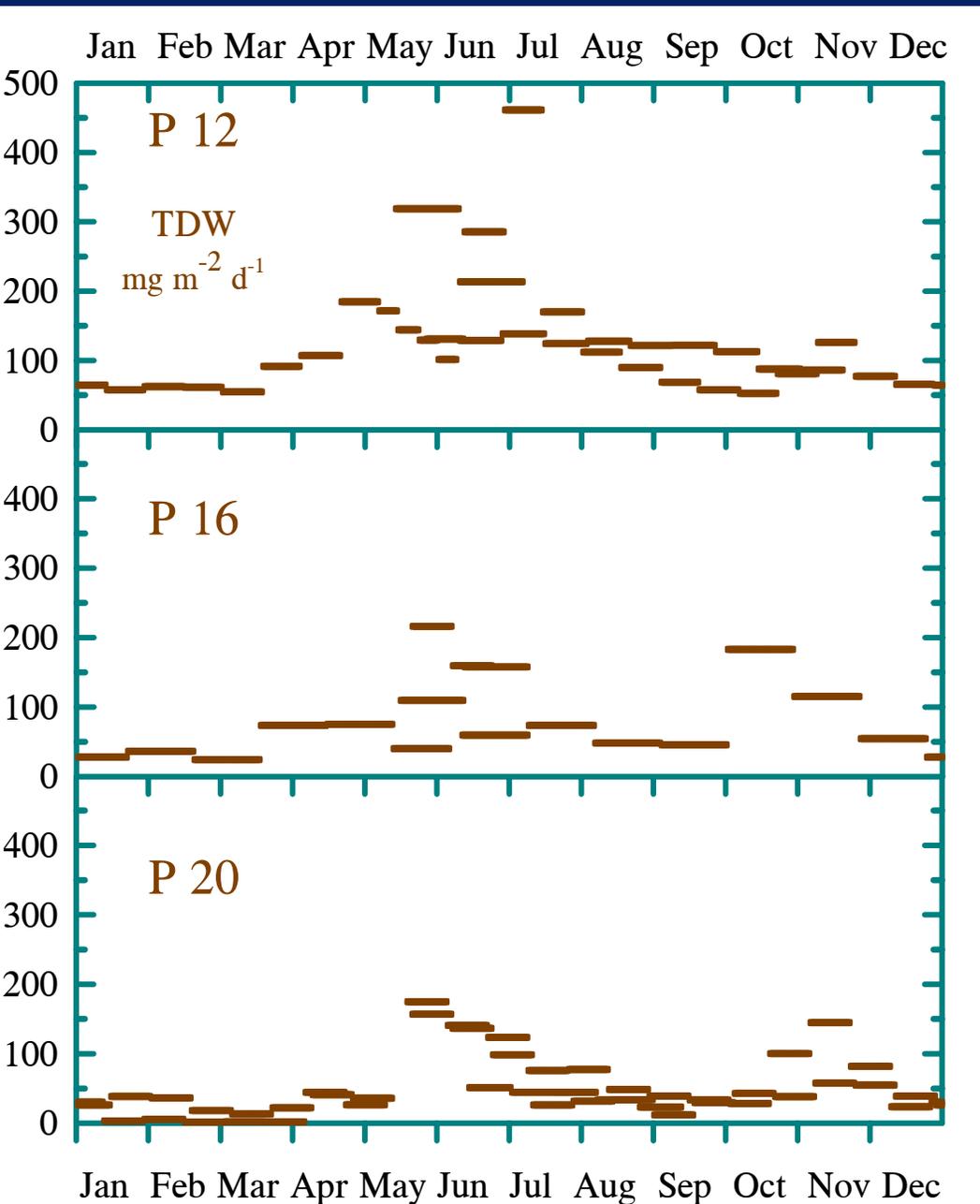


fluxes were recorded sporadically along line P between 1995 and 1998



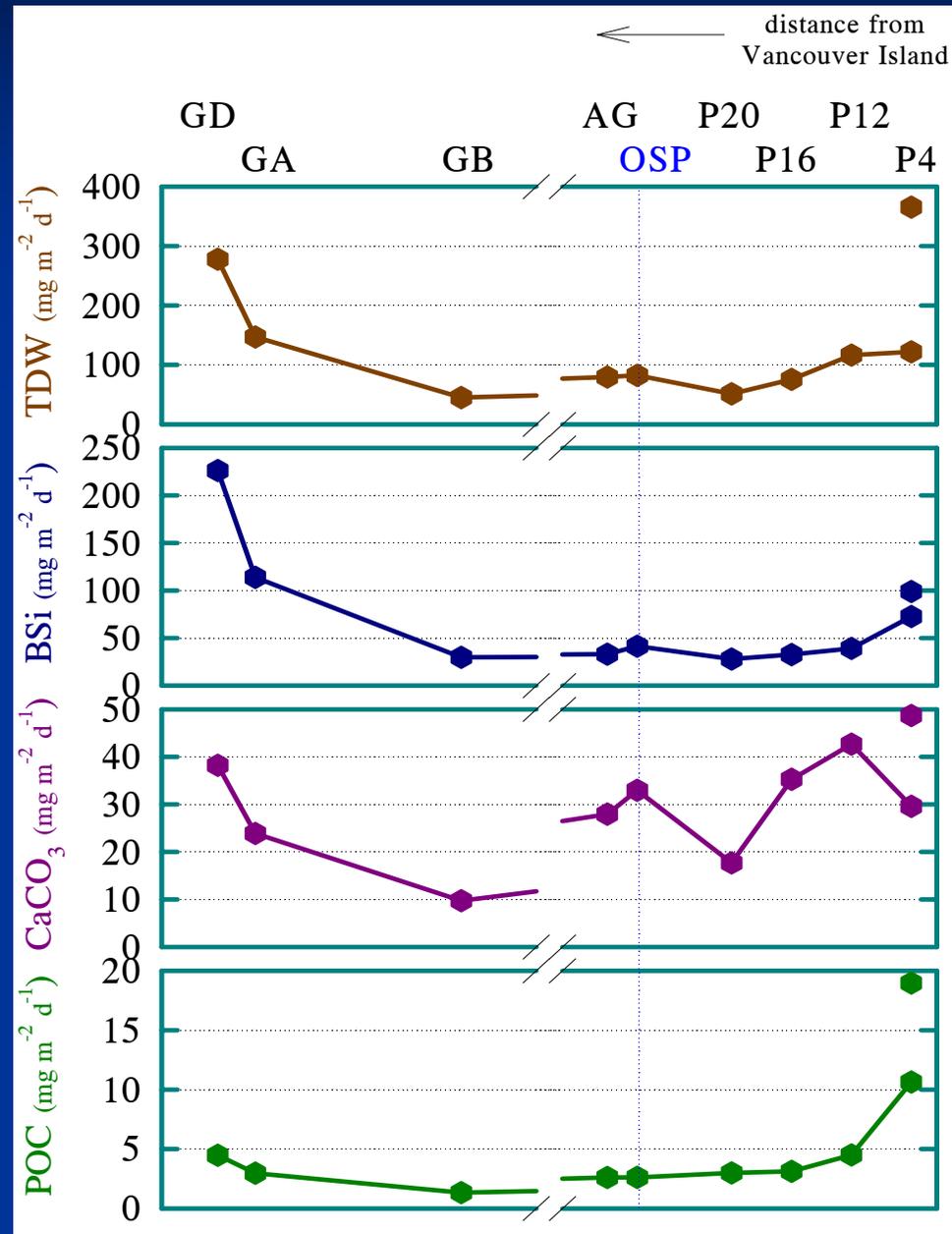
clogs!!!

but sorted by Julian day...



Summary of fluxes in the North Pacific:

mean fluxes measured at deepest traps at each station

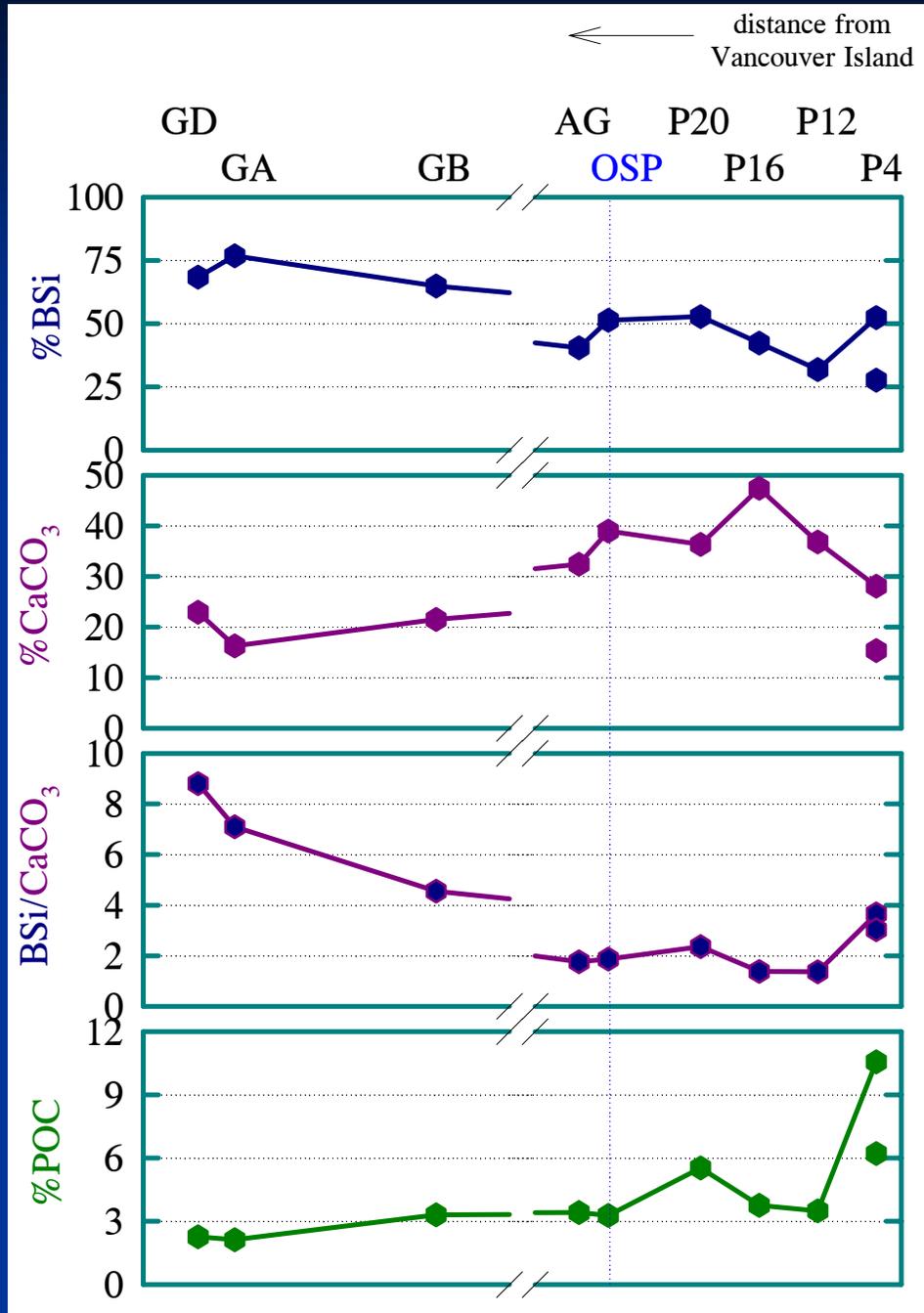


total flux at La Perouse Bank

“downward” flux

maximum coccolithophorid biomass along Line P periodically occurs at P12 (M Lipsen, pers. comm.).

highest BSi
contributions occur
In the Western
Pacific

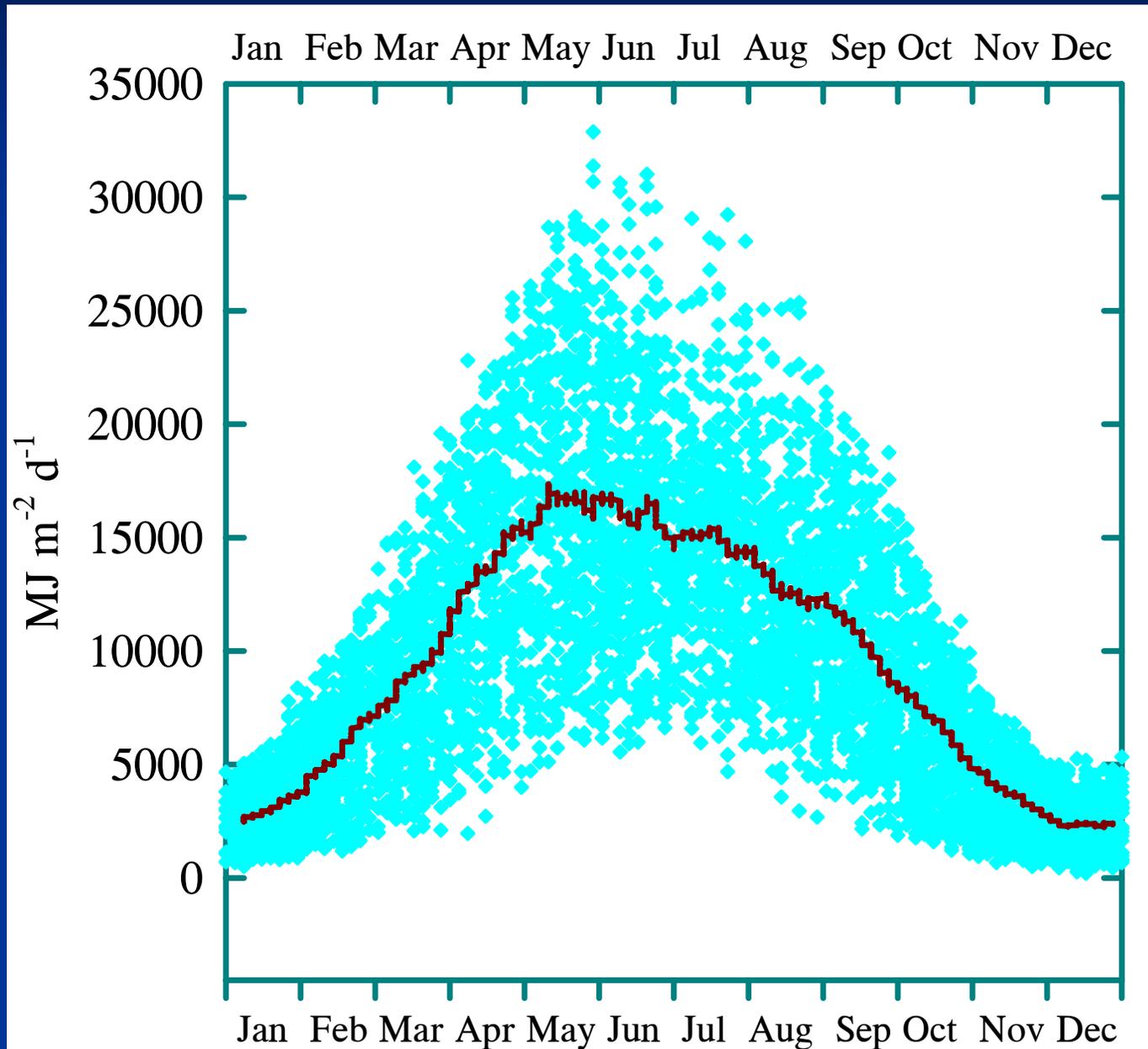


highest carbonate
contribution occurs
from P12 into the
Alaska Gyre
Foram?
continuous plankton
Recorder might help

conclusions

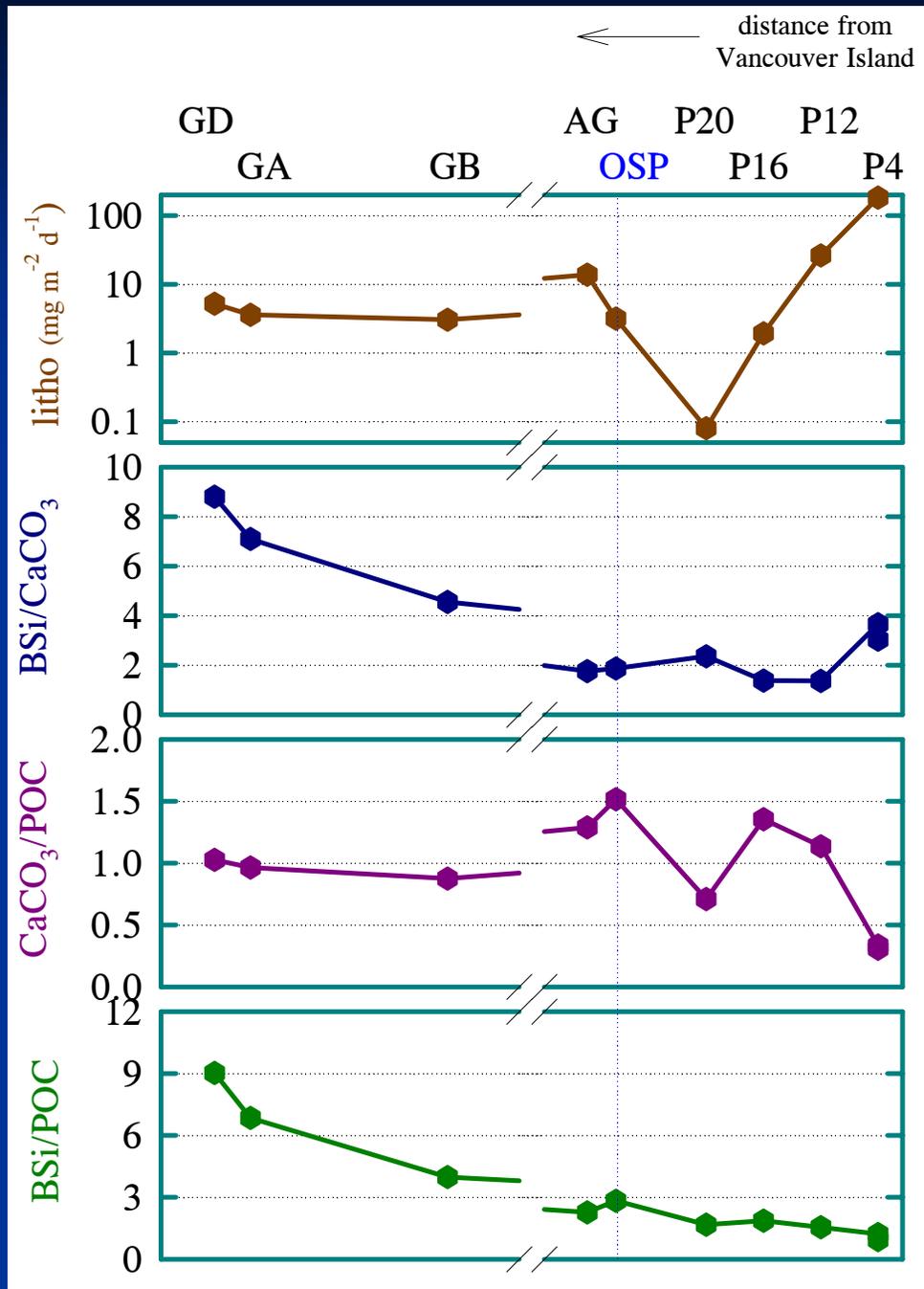
- Fluxes in the Alaska Gyre at 55°N were similar in magnitude and composition to fluxes at OSP.
- Fluxes at 3 stations west of the dateline were composed almost entirely of BSi. Total fluxes were very low near the dateline but in the NW, ~3x higher than in the eastern North Pacific.
- Fluxes along Line P decreased with distance from the coast. The “downward” flux at La Perouse Bank shows spring and summer peaks similar to those at OSP. Highest carbonate fluxes throughout the North Pacific were measured at P12 and P16.

light at OSP: 1960-1980

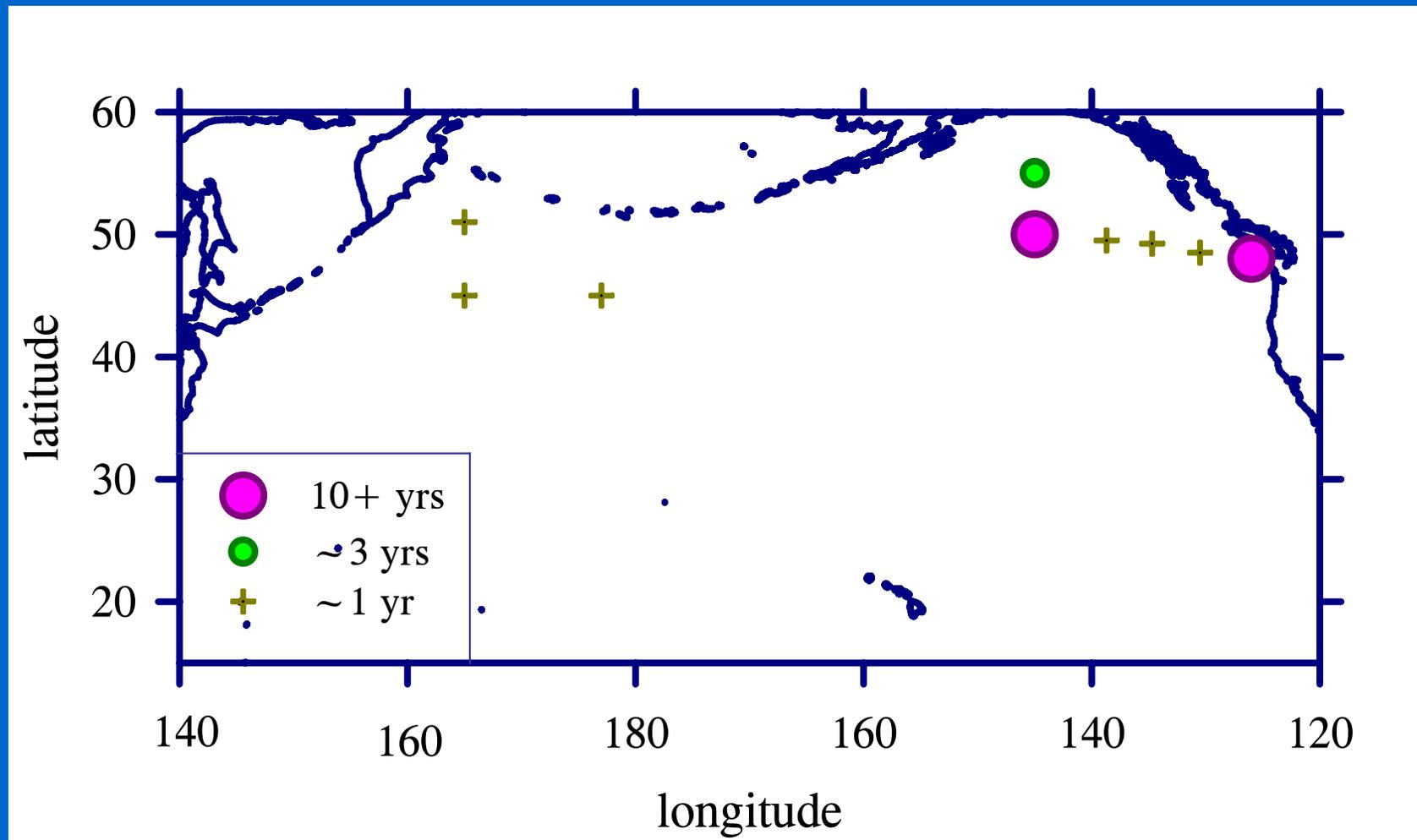


This slide will contain...

- Four figures (one for each constituent), exactly like upper right of the previous four slides, but with 200 m fluxes along with 1000 m and 3800 m. Point will be that the same signal exists at 200 m. A lot less attenuation than expected for the 200 m – 1000 m depth interval. There are some analytical issues we're dealing with, so we haven't yet pushed the 200 m time series through the calculations.



La Perouse Bank: 1986 to present



Southern Oscillation
Index

Pacific Decadal
Oscillation

BSi/CaCO₃ flux
anomaly

POC/CaCO₃ flux
anomaly

BSi/POC flux
anomaly

